Rivers and **Dams**

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USBR Short Course

1. Watersheds and Climate

- 2. Sedimentation Engineering
- **3. Rivers and Dams**
- 4. River Environment



Rivers and Dams

River Equilibrium
 Aggradation
 Degradation below Dams
 Case Study Gupo Bridge
 Case Study Dam Break



1a. Manning n



Rhine River flood in 1998





Primary dune height vs discharge



Manning n vs discharge



1b. Downstream Hydraulic Geometry



Downstream Hydraulic Geometry



Julien-Wargadalam (J-W) Equations

When the Manning-Strickler approximation is applicable, i.e. m = 1/6, a simplified form of Eqs. (10.19) is obtained in SI as

$$h \approx 0.133 Q^{0.4} \tau_*^{-0.2} \qquad (10.20a) \blacklozenge$$

$$W \approx 0.512 Q^{0.53} d_s^{-0.33} \tau_*^{-0.27} \qquad (10.20b) \blacklozenge$$

$$V \approx 14.7 Q^{0.07} d_s^{0.33} \tau_*^{0.47} \qquad (10.20c) \blacklozenge$$

$$S \approx 12.4 Q^{-0.4} d_s \tau_*^{1.2} \qquad (10.20d) \blacklozenge$$

The hydraulic geometry of stable channels is obtained from Eqs. (10.20) when $\tau * \cong 0.047$. Higher sediment transport rates require higher velocity and slope, and reduced width and depth.

Bankfull width and depth



from Julien and Wargadalam (ASCE-JHE, 1996)

Bankfull velocity and slope



from Julien and Wargadalam (ASCE-JHE, 1996)

1c. Meandering



Sediment Transport in Sharp Bends



Fig. 3. Relation between SF and θ_1/ϕ (M/N = 1).

Laboratory experiments show that fine sand can deposit where coarse sand cannot, i.e. point bars



from Kawai and Julien (JHR-IAHR, 1996)

Sediment Transport in Sharp Bends



Field measurements in the sharp bends of the Fall River, Colorado demonstrate that particles of different sizes move in different directions.



Fig. 5 Center of mass curves for three bedload size fractions

from Julien and Anthony (JHR-IAHR, 2002)

Meandering Simulations

T=0.5 hrT=1.0 hrT=1.5 hrs T=2.0 hrs 1111 T=2.5 hrs T=3.0 hrs

Model from Duan and Julien (ESPL, 2005)

Initial Conditions

- sine-generated
- deflection angle 30°
- discharge 2.1 *l*/s
- width 0.4 m
- length 13.2 m
- sediment size 0.45 mm

Meandering Evolution

Example starting from a straight channel on the Rio Puerco, New Mexico





Rio Puerco, New Mexico



Lateral Migration in a Meandering Channel



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2a. Meandering to Braiding (sediment overload)



Natural Chute Cutoffs

Often in

to an

load

•

response increase in sediment

> Chute cutoffs on Williams River, AK (Photo by N.D. Smith)



Oxbow Lake



Riverbed rising forces river overbank

2b. Channel width variability











From Julien River Mechanics CUP 2018



2c. Width-slope trade-offs





Relationship between channel width and sediment transport



Wider reaches are steeper!



From Leon et al. ASCE-JHE 135(4), 2009

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3a. Degradation Problems














3b. Braiding to meandering (sediment starved)





Planform geometry 1935 1972



1992



0

Lateral Adjustments



Cross section CO-18



Active channel width



Changes in active channel width Rio Grande, NM (after Richard et al., 2005)



Exponential Model Results



Hydraulic Geometry Equations (Julien & Wargadalam 1995)



3c. River Problems in Estuaries



Sand and Gravel Mining







Longitudinal Flood Profile for Sg Muda (Q=1340m³/s)



Received Receiving

Real South South South South South South

Guno bridge

Circusancity // K



•from Ji et al. ASCE-JHE, Nov. 2011

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Case-Study: Gupo Bridge during Typhoon Maemi in 2003









Alternative plan III





Figure 14. Gupo and Subway Bridge Piers before and after retrofitting construction

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How does the flood wave propagate downstream?

21 Mile Dam, Elko County, NV, Feb 8, 2017

25/01/2019 12:28:2

AO V

M1 - Barragem



The Fundão Dam Collapse

Bento Rodrigues Town, 6 November 2015



Gesteira Town 6 November 2015





Candonga Dam, 6 November 2015 C ≈ 700,000 mg/l Coast, 21 November 2015 C ≈1,500 mg/l



Doce River at Governador Valadares City 11 November 2015, C ≈30,000 mg/l



Hydrographs

Observed hydrographs in Doce River after the Fundão Dam break (ANA, 2015).



from Marcos Palu



Floodwave Propagation Modeling













from Marcos Palu and Julien ASCE-JHE (2020)

Sediment Concentration Measurements

Observed suspended sediment concentration (CPRM & ANA, 2015)



Sediment Routing NEW Development!

The one-dimensional advection-dispersion equation with settling is applied on the evaluation of transport of suspended load in open channels (Fischer et al., 1979; Julien, 2010).

$$\frac{\partial C}{\partial t} + U \frac{\partial C}{\partial x} = K_d \frac{\partial^2 C}{\partial x^2} - kC$$

C is the concentration: U is the flow averaged velocity; *K_d* is the longitudinal dispersion coefficient;

 $k = \omega/h$ is the settling rate.

Analytical solution for a constant spill of finite duration τ is:

$$C(x,t) = \frac{C_0}{2} \begin{cases} e^{\frac{Ux}{2K_d}(1-\Gamma)} \left[erfc\left(\frac{x-Ut\Gamma}{2\sqrt{K_d}t}\right) - erfc\left(\frac{x-U(t-\tau)\Gamma}{2\sqrt{K_d}(t-\tau)}\right) \right] \\ + e^{\frac{Ux}{2K_d}(1+\Gamma)} \left[erfc\left(\frac{x+Ut\Gamma}{2\sqrt{K_d}t}\right) - erfc\left(\frac{x+U(t-\tau)\Gamma}{2\sqrt{K_d}(t-\tau)}\right) \right] \end{cases}$$

$$Where: \ \Gamma = \sqrt{1+4\eta} \text{ and } \eta = \frac{kK_d}{u^2}$$
The complementary error function *erfc* is equal to:
$$erfc(b) = 1 - erf(b) = 1 - \frac{2}{c} \int_{-\infty}^{b} e^{-\beta^2} d\beta$$

$$erfc(b) = 1 - erf(b) = 1 - \frac{2}{\sqrt{\pi}} \int_{0}^{b} e^{-\beta}$$

erfc.precise in Xcel
Sediment Concentration Modeling



from Palu and Julien ASCE=JHE (2018)

Summary and Conclusions

1. River Equilibrium

Rivers can reach equilibrium after several years.

- 2. Aggradation
- Aggradation forces out-of-banks rivers and braiding.
- 3. Degradation below Dams

Degradation causes incision and narrowing with possible impact on structures.

4. Case-study Gupo Bridge

Retrofitting based on stability, not equilibrium.

5. Case-study Dam Break

Numerical modeling of flood and sediment waves.





Second Edition

Thank You!

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