



Sedimentation Engineering

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Colorado State University

USBR Lectures – Part II

Denver, Colorado

January 17, 2024

USBR Short Course

1. Watersheds and Climate
2. **Sedimentation Engineering**
3. Rivers and Dams
4. River Environment



Sedimentation Engineering

- 1. Incipient motion**
- 2. SEMEP and SEMEPP**
- 3. Mudflows and Debris Flows**



Shields Parameter

- Shields only defines the ratio of the shear force (no lift) to the particle submerged weight

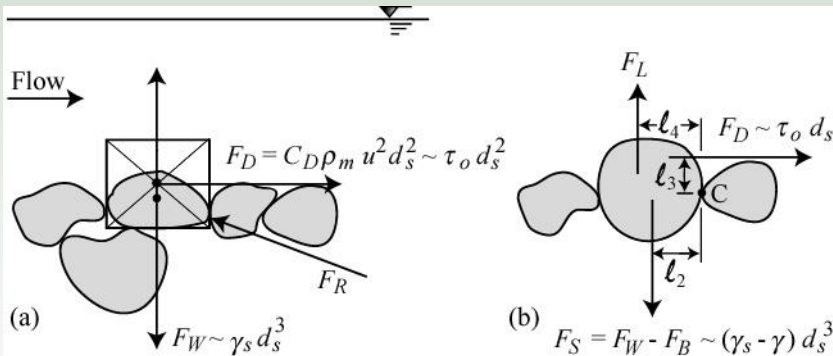


Figure 7.6. Force diagram under steady uniform flow

$$\tau_* = \frac{\tau_o}{(\gamma_s - \gamma)d_s} = \frac{\rho_m u_*^2}{(\gamma_s - \gamma)d_s}$$

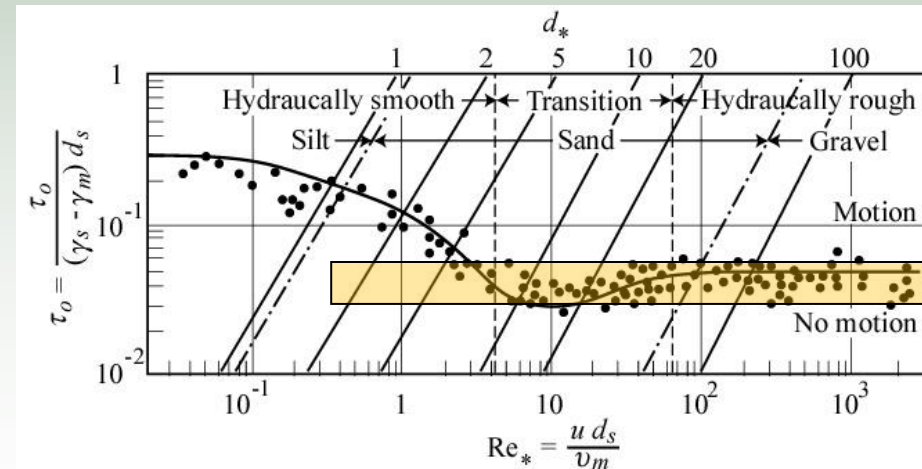
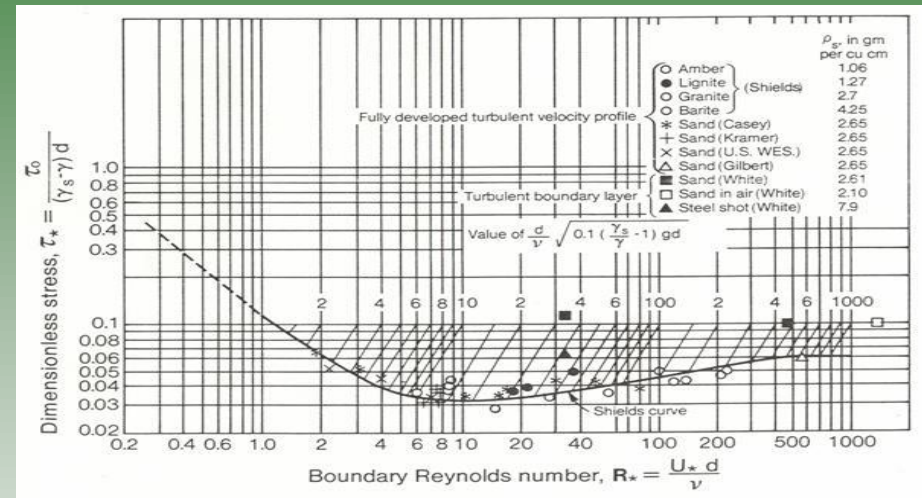
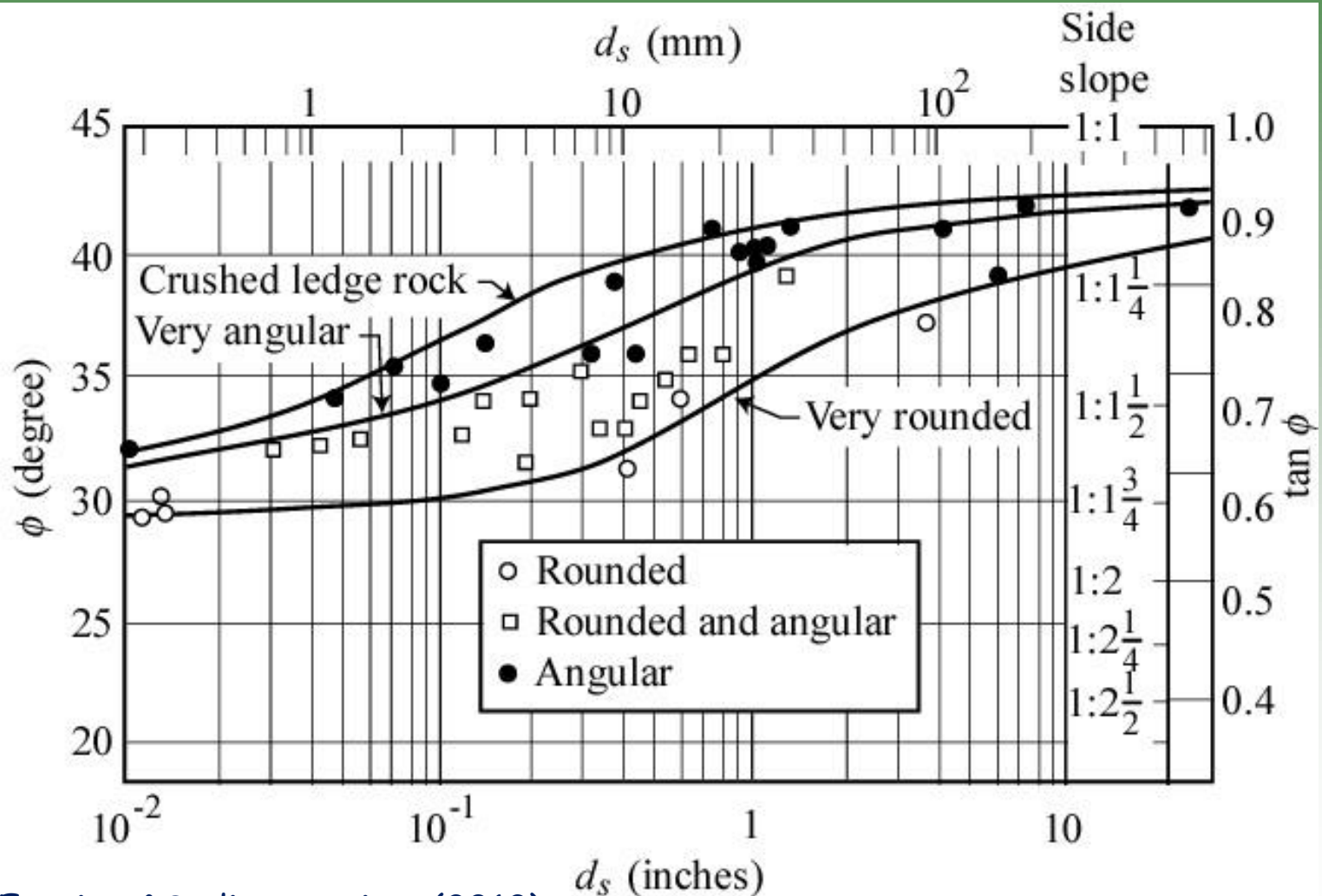


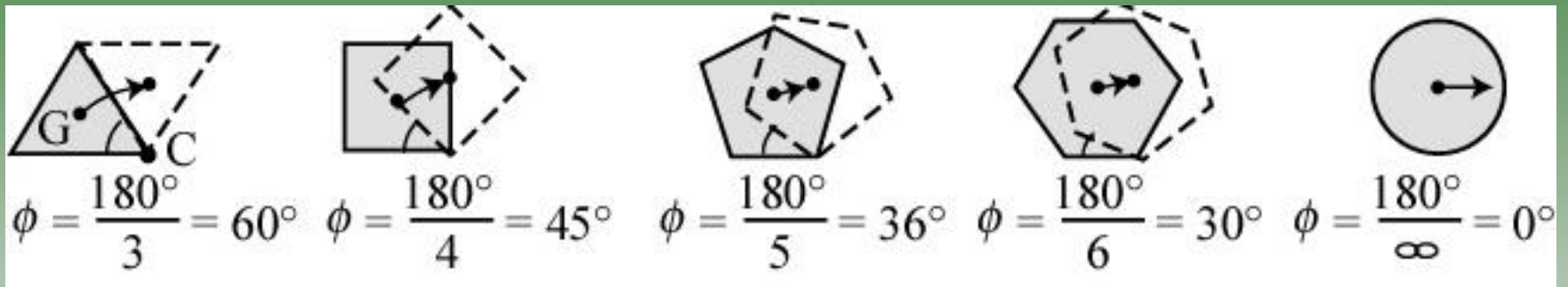
Figure 7.7. Shields diagram for granular material

Angle of Repose Granular Material

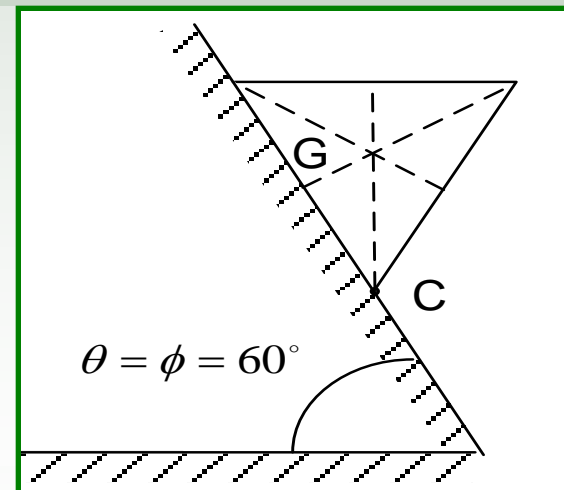


Angle of Repose

Effects of Angularity

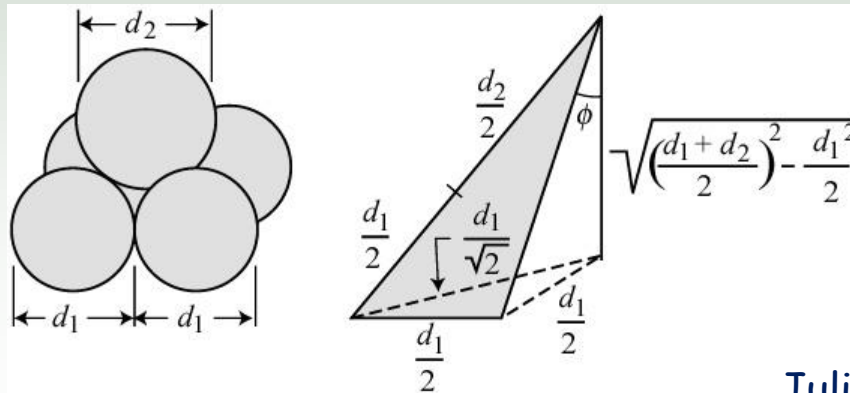
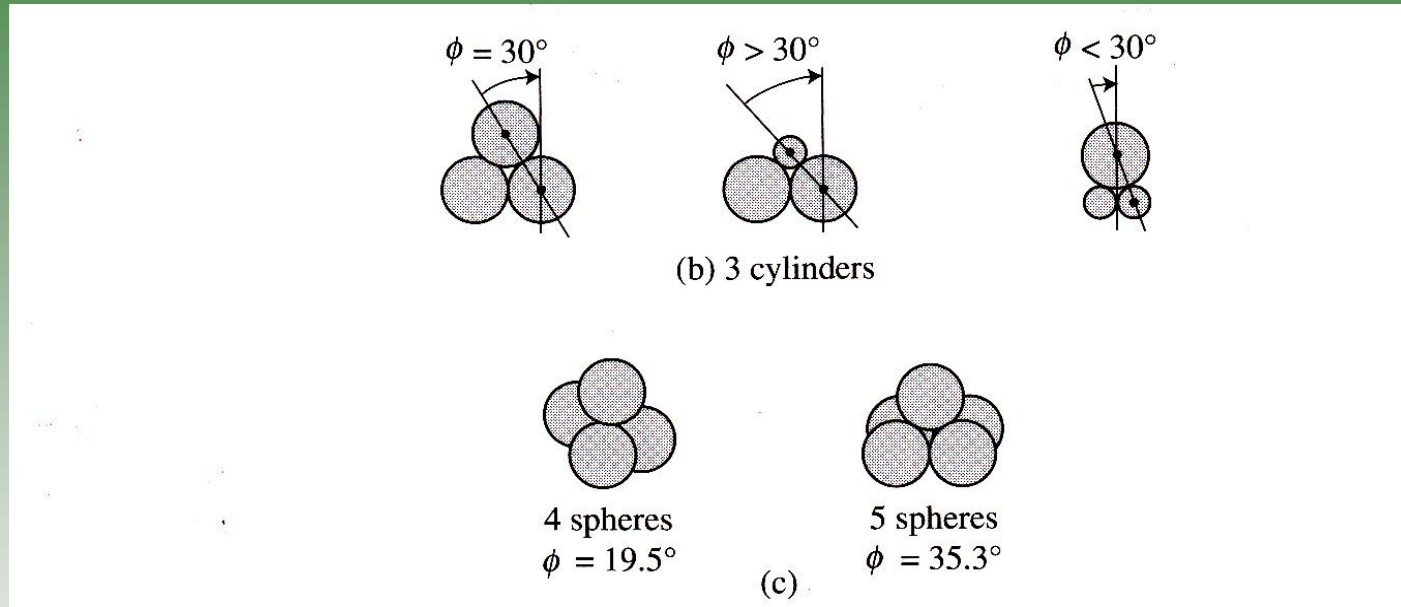


Motion occurs when the center of gravity (G), is above the point of contact (C).



Angle of Repose

Boundary and Particle Configuration



$$\tan \phi = \frac{d_1}{\sqrt{(d_1 + d_2)^2 - 2d_1^2}}$$

Julien Erosion & Sedimentation (2012)

Figure 7.3. Angle of repose for spheres of different diameter

Angle of Repose Boundary and Particle Size

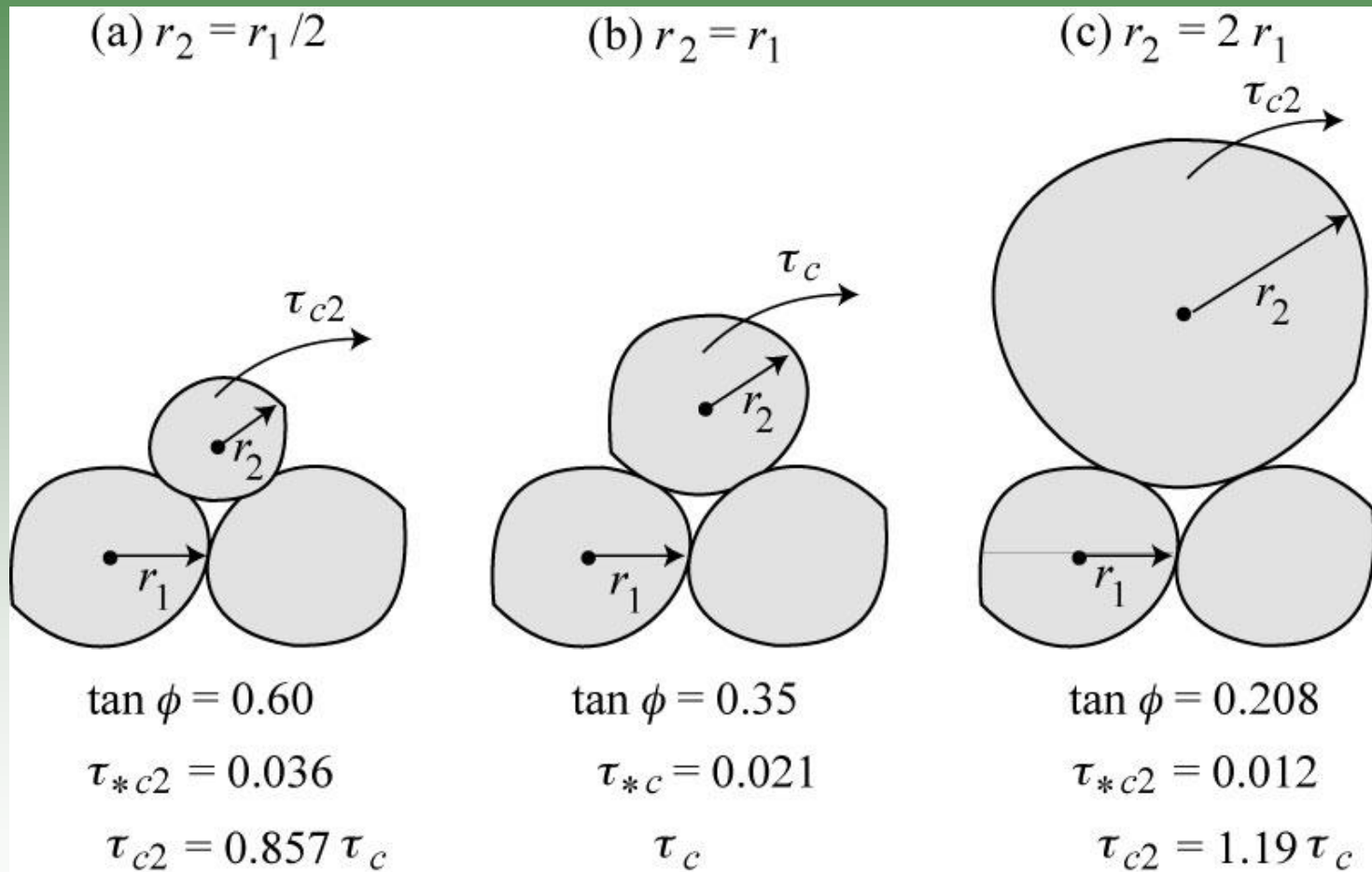
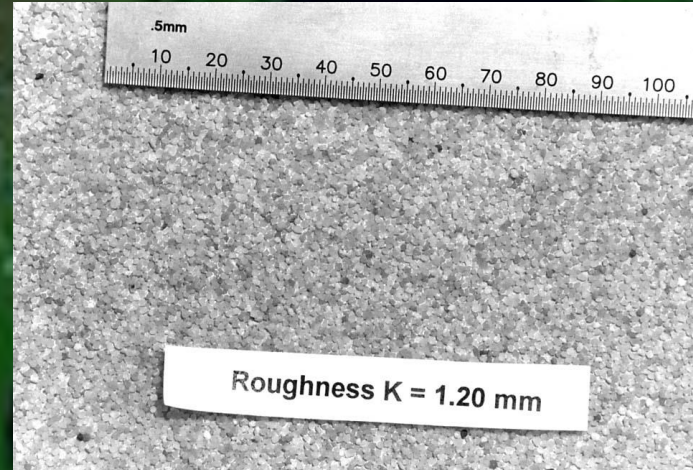
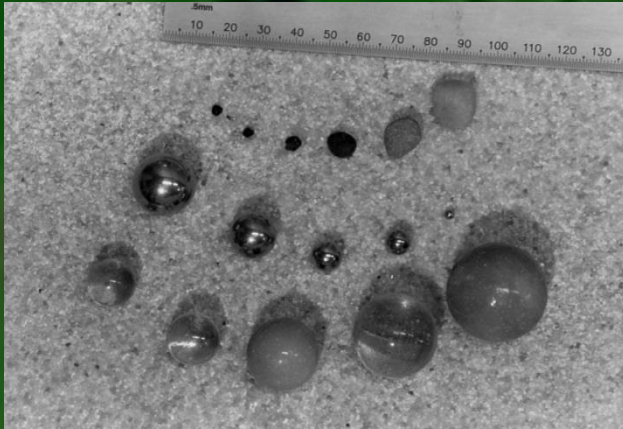
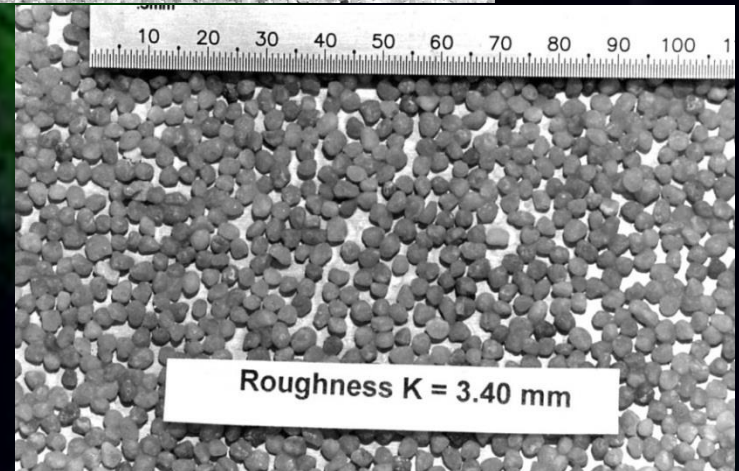


Figure 7.12. Relative stability for different particle sizes

Laboratory experiments at CSU



~10,000 particle velocity measurements



Bounvilay and Julien, ASCE-JHE (2013)

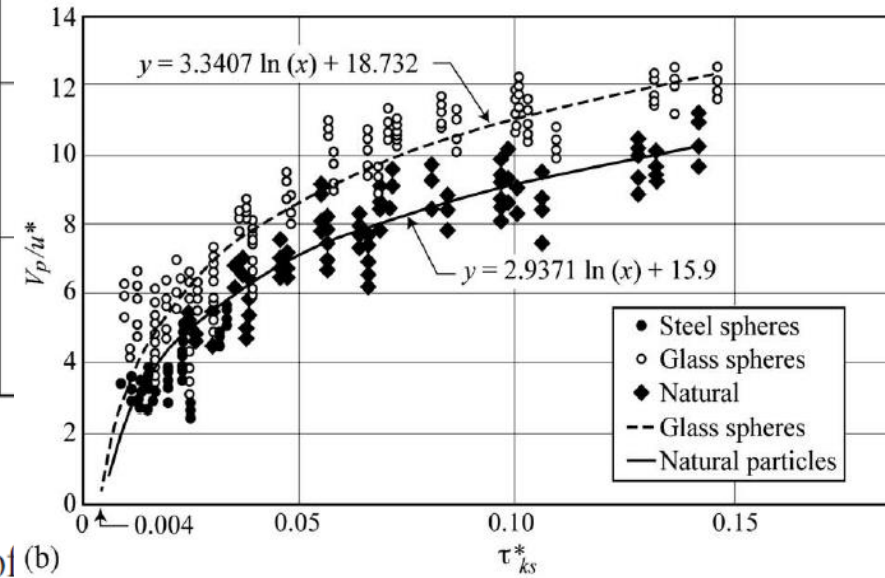
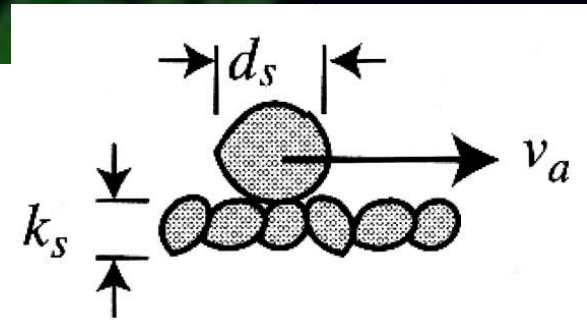
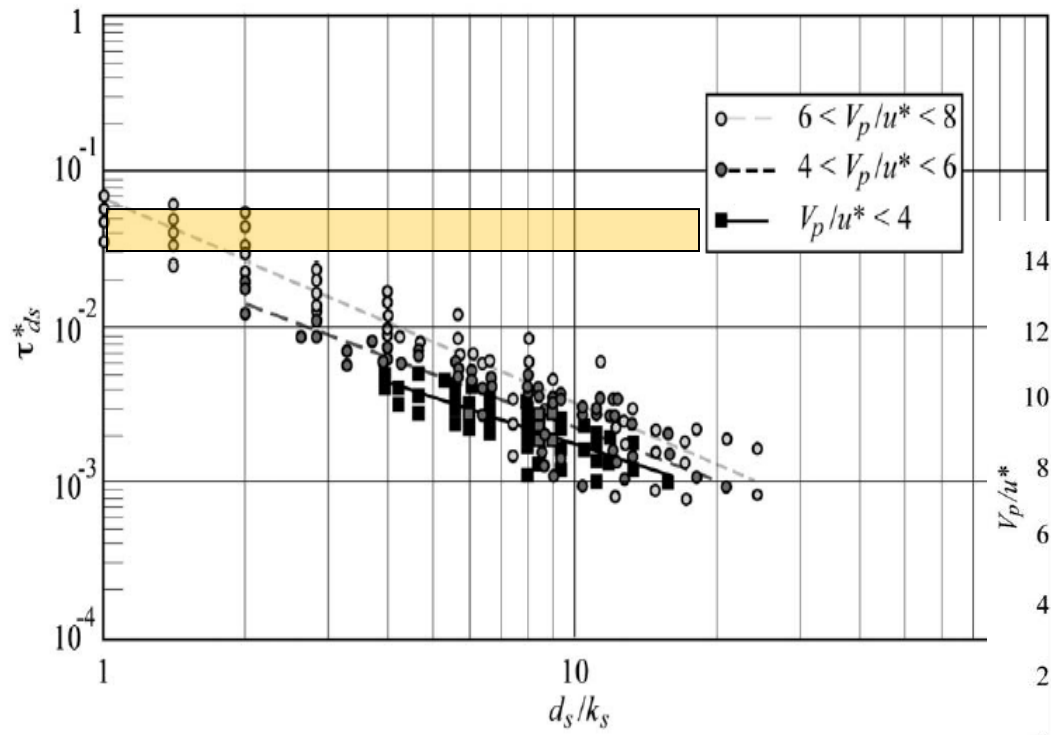


Fig. 9. Shields parameter $\tau_{d_s}^*$ vs d_s/k_s for different values of (b) V_p/u_*

Bounvilay and Julien
ASCE-JHE (2013)

Lets remember:
Incipient motion depends on the sum of moments, not forces.
Particles move at Shields values $\ll 0.03$ (in shaded box)

Sedimentation Engineering

1. Incipient motion
2. **SEMEP and SEMEPP**
3. Mudflows and Debris Flows

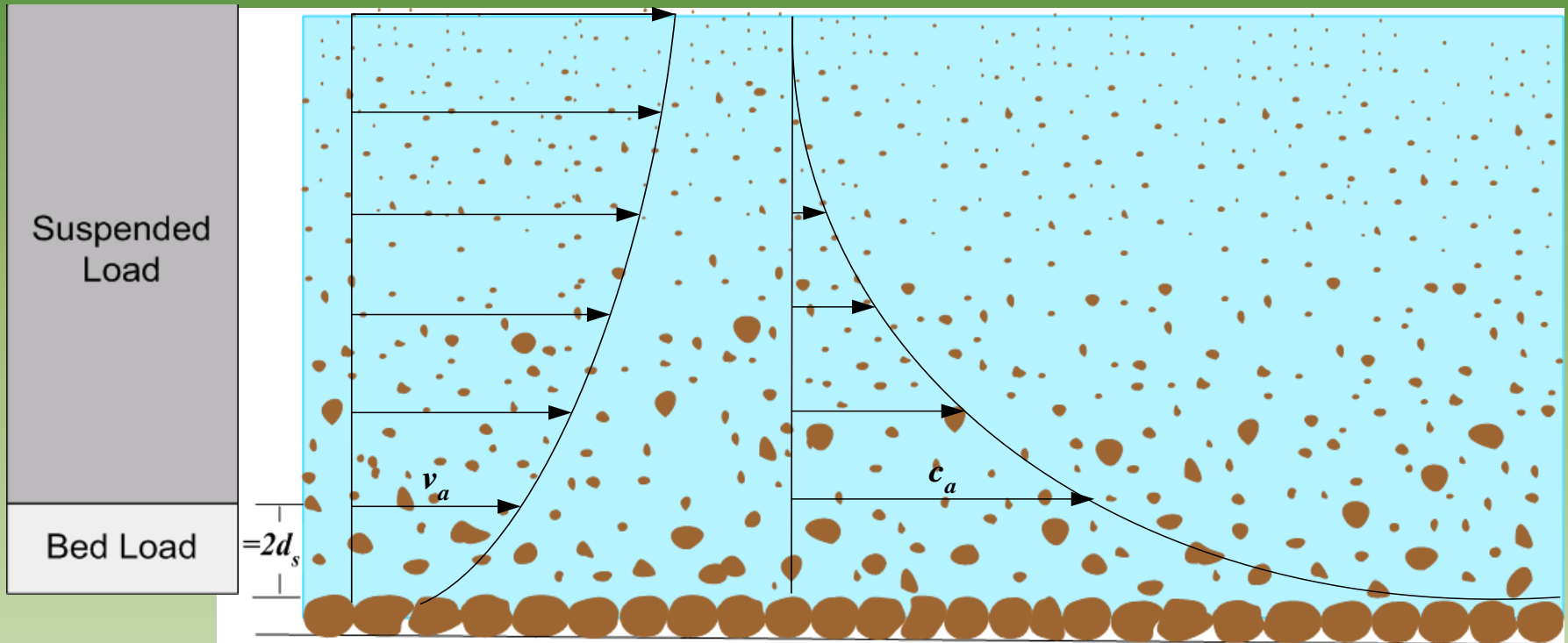


Sediment Transport Mechanism

Einstein -> bottom up
 Modified Einstein (MEP) -> top down

Velocity Profile

Concentration Profile

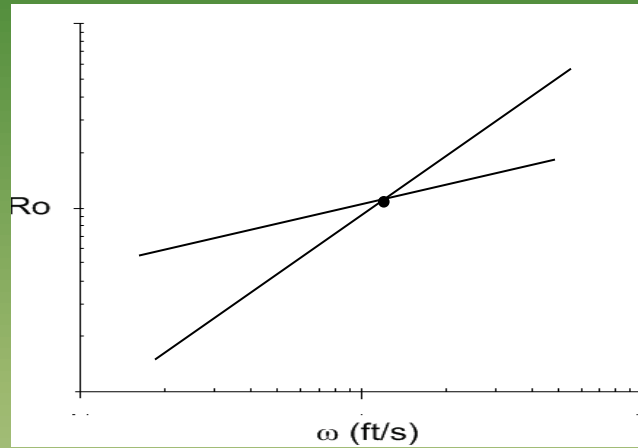
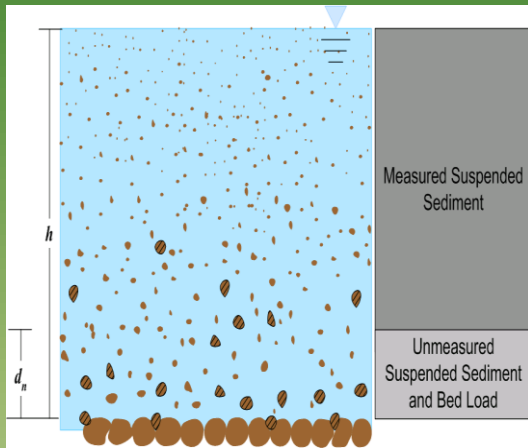


$$v = \frac{u_*}{\kappa} \ln \left(30.2 \frac{y}{k_s} \right)$$

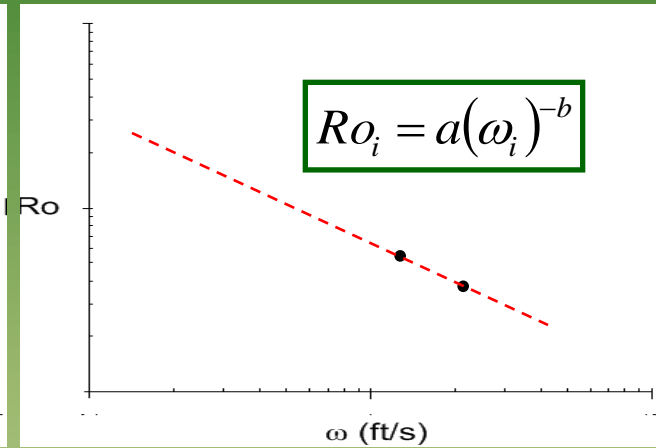
$$c = c_a \left(\frac{h-y}{y} \frac{a}{h-a} \right)^{Ro}$$

$$Ro = \frac{\omega}{\kappa u_*}$$

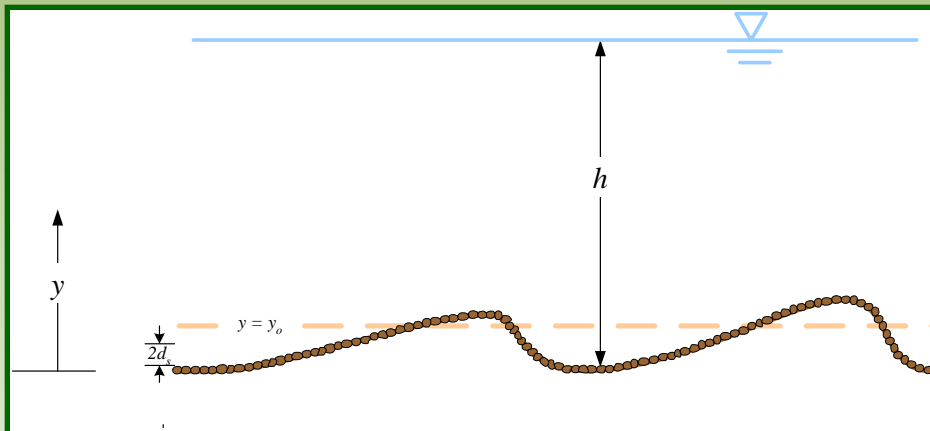
Problems with the 'traditional' Modified Einstein Procedure



CASE 1 – Not Enough Overlapping Bins



CASE 2 – Negative exponent



CASE 3 – Calculate Total Load is less than Suspended Load

Series Expansion Modified Einstein Procedure SEMEP

Efficient Algorithm for Computing Einstein Integrals

Junke Guo¹ and Pierre Y. Julien²

Einstein Algorithm Calculator:

This calculator was made using equations in the following reference:

Junke Guo and Pierre Y. Julien (2004). "Efficient Algorithm for Computing Einstein Integrals." J. Hydraul. Eng., 130(12), 1198-1201.

DOI: 10.1061/(ASCE)0733-9429(2004)130:12(1198)

Procedure:

The cells with red are inputs while the cells with blue are outputs.

To calculate, place mouse on user input cell, type in the desired inputs and press the enter key.

To change the inputs, simply select cell again and type in desired input and press the enter key.

To view the equations used in the calculations select the Einstein Algorithm Calculations tab below.

Also included are graphical solutions for $J_1(z)$ & $J_2(z)$ using E-values of 0.1 and .00001.

Non-Integer Algorithm Analysis

		Value Inputted	Value Used In Analysis
Input:	z=	1	1
	E=	0.00001	0.00001

Ouput:	$J_1(z)=$		$J_2(z)=$
	$J_1(0)=$		$J_2(0)=$
	$J_1(1)$	1.0515E+01	$J_2(1)$
			6.5935E+01

Primary Mode of Transport from SEMEP

Ratio of the suspended to total load q_s/q_t

- Calculate the Suspended Load

$$q_s = \int_{2ds}^h cv dy$$

$$q_s = \int_{2ds}^h c_a \left(\frac{h-y}{h} \frac{a}{h-a} \right)^{Ro} \frac{u_*}{\kappa} \ln \left(\frac{30y}{d_s} \right) dy$$

$$c_a = \frac{q_b}{v_a a}$$

$$v_a = 11.6u_*$$

$$q_s = 0.216q_b \frac{E^{Ro-1}}{(1-E)^{Ro}} \left\{ \ln \left(\frac{30h}{d_s} \right) J_1 + J_2 \right\}$$

- Calculate the Total Load

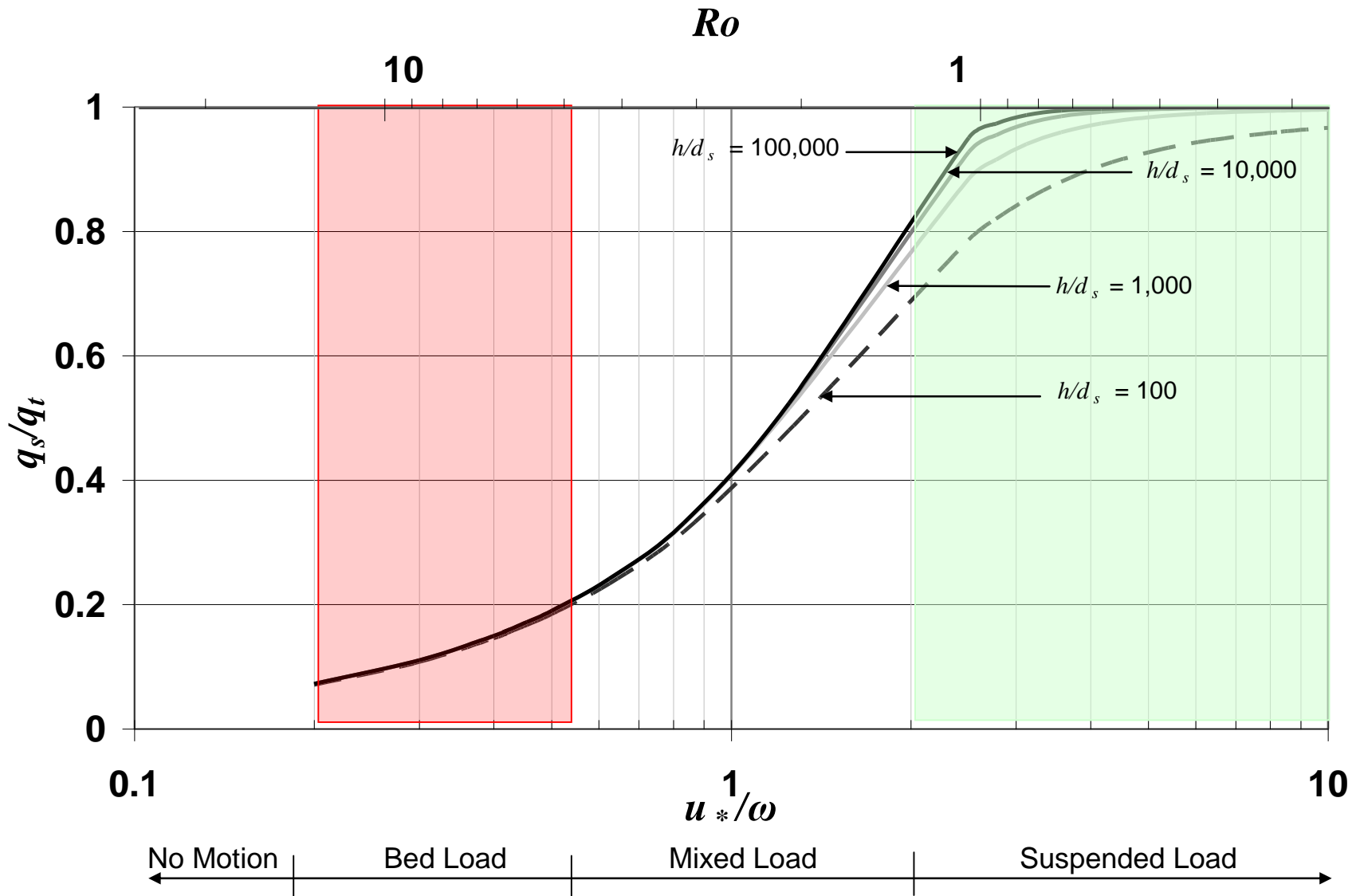
$$q_t = q_b + q_s$$

$$q_t = q_b + 0.216q_b \frac{E^{Ro-1}}{(1-E)^{Ro}} \left\{ \ln \left(\frac{30h}{d_s} \right) J_1 + J_2 \right\}$$

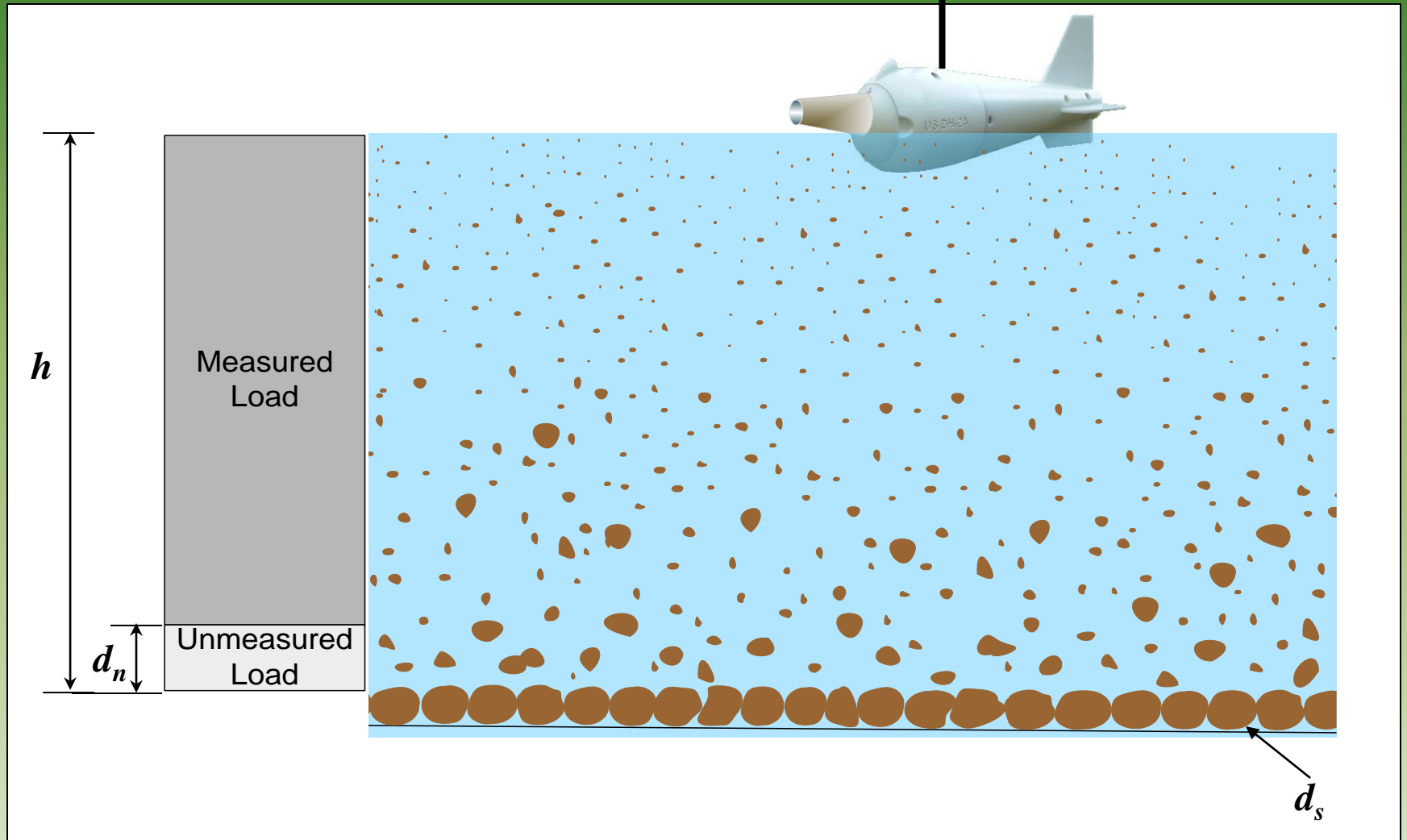
- Ratio of q_s/q_t

$$\frac{q_s}{q_t} = \frac{0.216 \frac{E^{Ro-1}}{(1-E)^{Ro}} \left\{ \ln \left(\frac{30h}{d_s} \right) J_1 + J_2 \right\}}{1 + 0.216 \frac{E^{Ro-1}}{(1-E)^{Ro}} \left\{ \ln \left(\frac{30h}{d_s} \right) J_1 + J_2 \right\}}$$

Primary Mode of Transport



Ratio of Measured to Total Load from SEMEP



Total Sediment Load from SEMEP Using Depth-Integrated Concentration Measurements

Seema C. Shah-Fairbank, M.ASCE¹; Pierre Y. Julien, M.ASCE²; and Drew C. Baird, M.ASCE³

Abstract: This study improves total sediment load calculations on the basis of depth-integrated sediment concentration measurements for channels with significant sediment transport in suspension. The series expansion of the modified Einstein procedure (SEMEP) removes most of the empiricism found in the existing modified Einstein procedures (MEP). SEMEP calculations require field measurements of flow discharge, depth-integrated suspended sediment (SS) concentration, and suspended particle sizes. SEMEP calculates the Rouse number, Ro , from the median particle size measured in suspension $d_{50,ss}$. On the basis of the sediment discharge measurements collected from 14 rivers, the accuracy of sediment discharge calculations depend on the ratio of the shear velocity u_* to the settling velocity ω . SEMEP performs accurately (error less than 25%) and without bias when $u_*/\omega > 5$. Calculations are also acceptable, but less accurate when u_*/ω is between two and five. Both SEMEP and MEP should not be used when the value of $u_*/\omega < 2$, and a simplified formulation on the basis of bed sediment discharge is recommended when $u_*/\omega < 2$. DOI: 10.1061/(ASCE)HY.1943-7900.0000466. © 2011 American Society of Civil Engineers.

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International Journal of Sediment Research 30 (2015) 1-12

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RESEARCH**
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Sediment load calculations from point measurements in sand-bed rivers

Seema C. SHAH-FAIRBANK¹ and Pierre Y. JULIEN²

Abstract

Point velocity and suspended sediment concentration measurements are used to calculate the total sediment discharge in sand-bed rivers. Calculations with the Series Expansion of the Modified Einstein Point Procedure (SEMEPP) depend on grain diameter d_s and settling velocity ω , flow depth h , shear velocity u_* , and sampling depth h_p . This procedure extends the applicability of the Modified Einstein Procedure (MEP) by using point sediment concentration and velocity measurements. This procedure is tested using the laboratory data from Coleman, and field measurements from the Enoree, Middle Rio Grande and Mississippi Rivers. Based on 801 point measurements over 124 verticals at flow depths ranging from 0.17 m to 33.5 m and sediment concentrations less than 0.1 kg L^{-1} , the accuracy of the calculations depends on u_*/ω and h_p/d_s . Point measurement techniques like SEMEPP are well-suited when $u_*/\omega > 5$ where at least 60% of the total sediment load is measured when 90% of the flow depth is sampled. The determination of sediment discharge from point measurements is most accurate in deep rivers when $h_p/d_s > 10,000$, and $u_*/\omega > 10$. Point measurements are not well-suited for shallow rivers and laboratory flumes where $h < 0.5 \text{ m}$ and when $u_*/\omega < 2$.

International Journal of Sediment Research 34

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Original Research

The ratio of measured to total sediment discharge

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Department of Civil and Environmental Engineering, Colorado State University, Fort Collins, CO 80523, USA

Zone of Applicability

$$q_m/q_t$$

- Calculate the Measured Load

$$q_m = \int_0^h cv dy$$

$$q_m = \int_0^h c_a \left(\frac{h-y}{h} \frac{a}{h-a} \right)^{Ro} \frac{u_*}{\kappa} \ln \left(\frac{30y}{d_s} \right) dy$$

$$c_a = \frac{q_b}{v_a a}$$

$$v_a = 11.6u_*$$

$$q_m = 0.216q_b \frac{E^{Ro-1}}{(1-E)^{Ro}} \left\{ \ln \left(\frac{30h}{d_s} \right) J_{1A} + J_{2A} \right\}$$

- Calculated Total Load

$$q_t = q_b + q_s$$

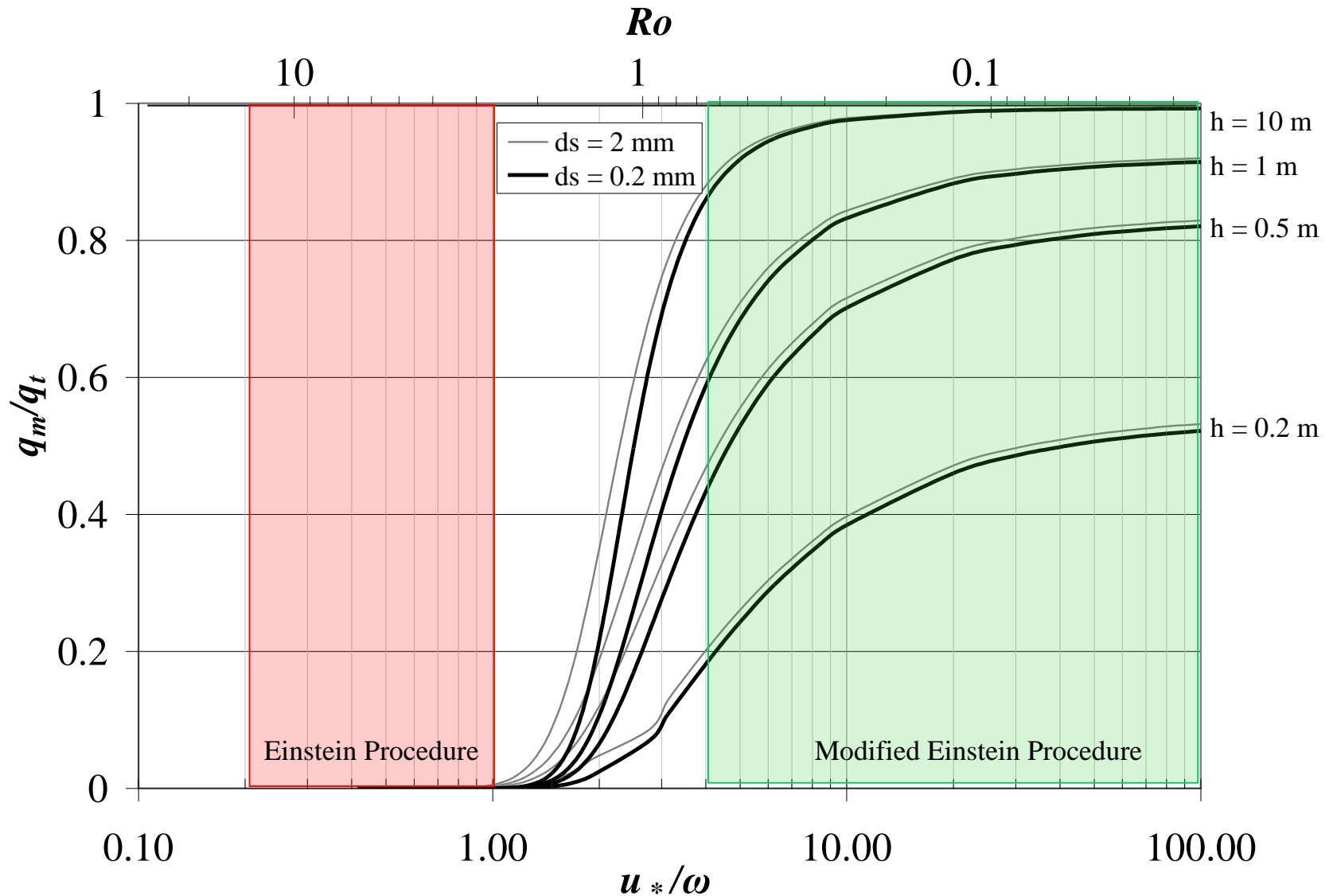
$$q_t = q_b + 0.216q_b \frac{E^{Ro-1}}{(1-E)^{Ro}} \left\{ \ln \left(\frac{30h}{d_s} \right) J_1 + J_2 \right\}$$

- Ratio of q_m/q_t

$$\frac{q_m}{q_t} = \frac{0.216 \frac{E^{Ro-1}}{(1-E)^{Ro}} \left\{ \ln \left(\frac{30h}{d_s} \right) J_{1A} + J_{2A} \right\}}{1 + 0.216 \frac{E^{Ro-1}}{(1-E)^{Ro}} \left\{ \ln \left(\frac{30h}{d_s} \right) J_1 + J_2 \right\}}$$

Measured to Total Sediment Load

from a standard suspended sediment sampler $d_n = 10$ cm



Field Testing by Seema Shah-Fairbank



Enoree River, SC



Susitna River, AK



Snake River, WA



Middle Rio Grande, NM



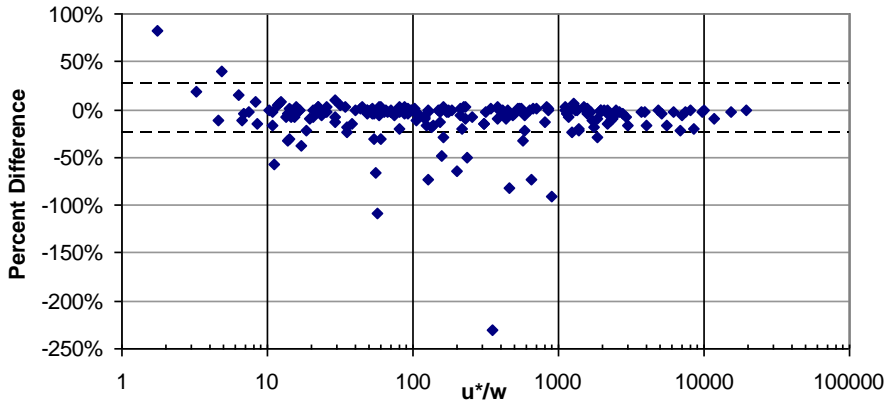
Clearwater River, ID



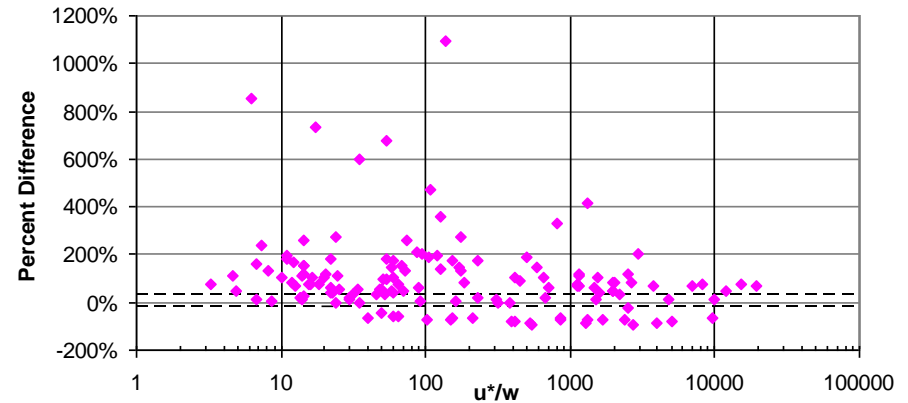
Mississippi River, MS

Results from Depth Integrated Sampler

Accuracy of Proposed Method



Accuracy of MEP



River	Hydraulic Data					Median Particle Size (mm)		Total Number of Samplers	Number of Samples with Calculated Total Load within 25% of measured Total Load	
	Q (cms)	V (m/s)	h (m)	W (m)	S	Bed	Suspension		Proposed Method	BORAMEP
Susitna River near Talkeetna Alaska	238 to 1,300	1.2 to 2.7	1.1 to 2.4	165 to 202	0.0011 to 0.0018	0.392 to 41.465	0.004 to 0.227	37	37	3
Chulitna River below Canyon near Talkeetna, Alaska	212 to 1,350	1.2 to 2.7	1.7 to 3.6	98.5 to 136	0.00039 to 0.011	0.744 to 11.273	0.006 to 0.500	43	33	0
Susitna River at Sunshine, Alaska	504 to 2,740	1.4 to 2.6	0.8 to 4.4	174 to 311	0.0012 to 0.0024	0.436 to 29.440	0.006 to 0.175	37	36	3
Snake River near Anatone, Washington	1180 to 37,770	1.9 to 3.3	3.9 to 5.8	169 to 197	0.000087 to 0.00124	0.412 to 43.077	0.0075 to 0.079	31	27	1
Toutle River at Tower Road near Silver Lake, Washington	29.2 to 592	1.2 to 3.1	0.46 to 2.3	20 to 70	0.0013 to 0.0055	0.583 to 20.766	0.025 to 0.220	19	16	9
North Fork Toutle River near Kid Valley, Washington	26.1 to 171	1.7 to 2.8	0.74 to 1.1	20.5 to 59	0.0032 to 0.0038	3.161 to 9.290	0.047 to 0.158	5	5	2
Clearwater River at Spalding, Idaho	677 to 2,740	1.2 to 3.4	4.1 to 5.5	128 to 146	0.000183 to 0.00056	0.390 to 27.636	0.005 to 0.291	35	32	1

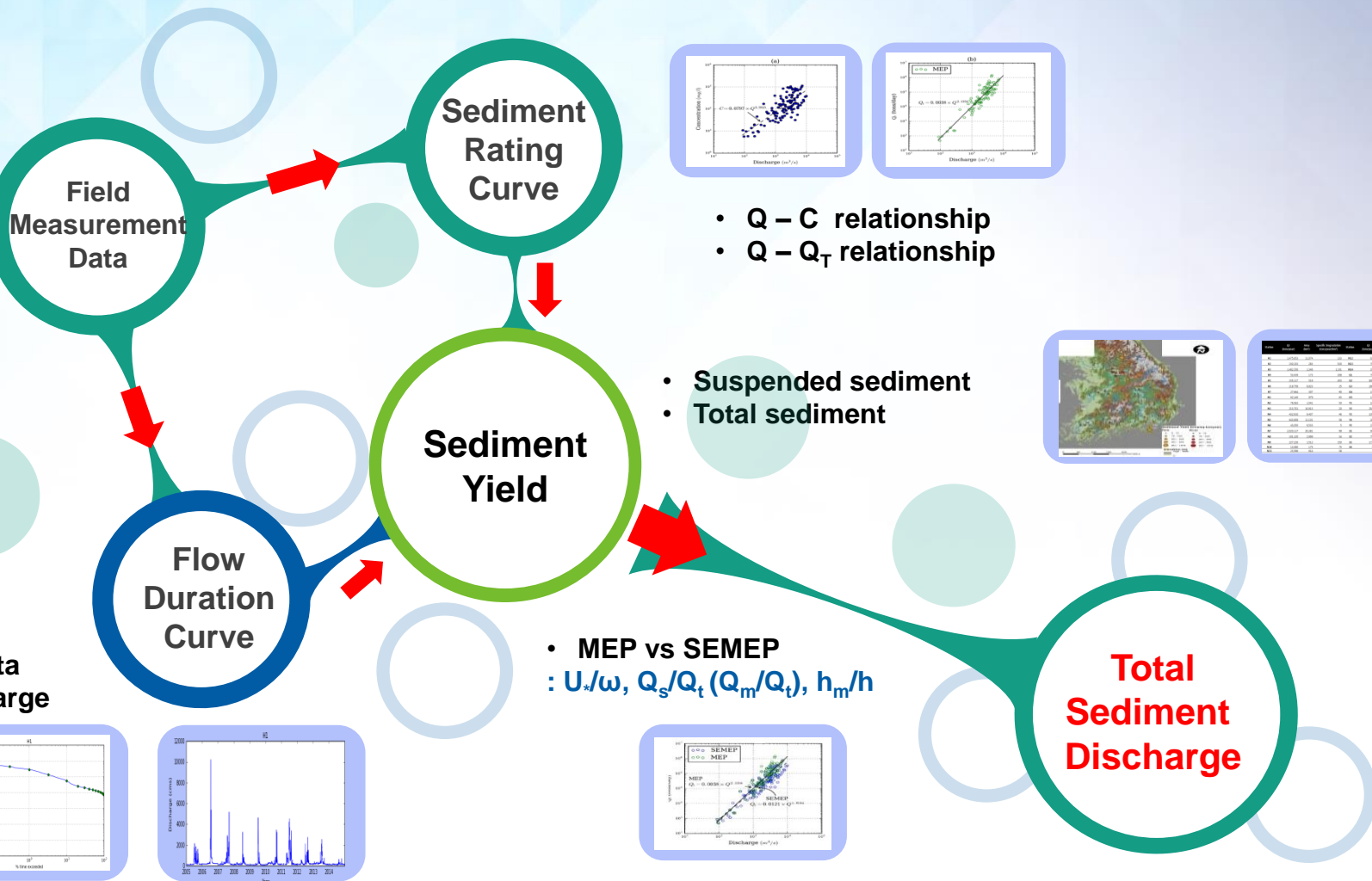
Sediment Calculations (All Korean Rivers)

Process for River Sediment

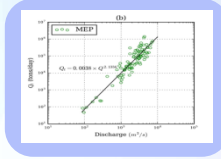
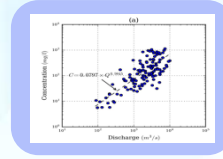
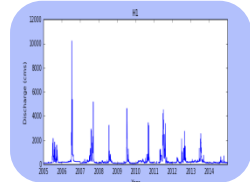
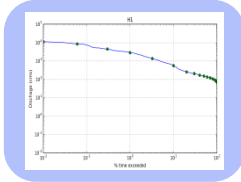


DATE	TIME	WATER LEVEL (m)	WIND SPEED (m/s)	WIND DIRECTION	TEMPERATURE (°C)	RELATIVE HUMIDITY (%)	PRECIPITATION (mm)	DISCHARGE (m³/s)	SEDIMENT CONCENTRATION (kg/m³)
2015-01-01	00:00	1.2	2.5	SE	10	80	0	100	0.5
2015-01-01	06:00	1.1	3.0	SE	11	75	0	120	0.6
2015-01-01	12:00	1.0	2.8	SE	12	70	0	150	0.7
2015-01-01	18:00	0.9	2.5	SE	11	75	0	100	0.5

(Korean team)

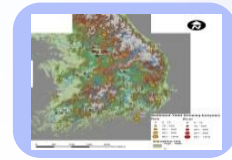


- 10 years data
- Daily discharge



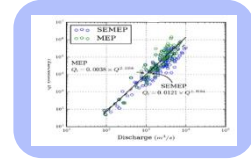
- Q – C relationship
- Q – Q_T relationship

- Suspended sediment
- Total sediment



DATE	TIME	WATER LEVEL (m)	WIND SPEED (m/s)	WIND DIRECTION	TEMPERATURE (°C)	RELATIVE HUMIDITY (%)	PRECIPITATION (mm)	DISCHARGE (m³/s)	SEDIMENT CONCENTRATION (kg/m³)
2015-01-01	00:00	1.2	2.5	SE	10	80	0	100	0.5

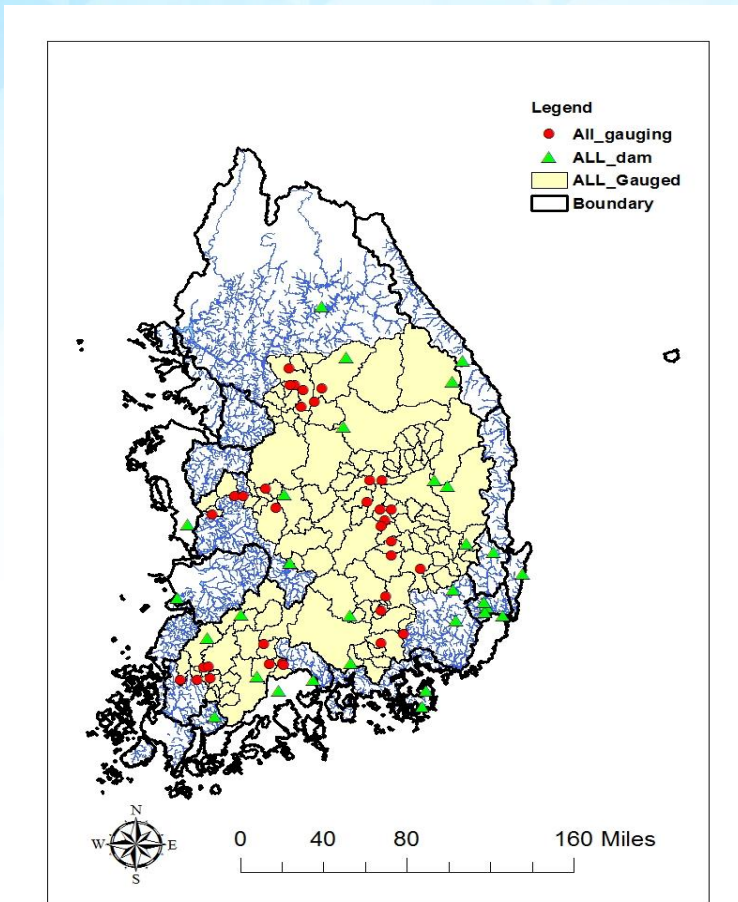
- MEP vs SEMEP
- : U_s/ω , Q_s/Q_t (Q_m/Q_t), h_m/h



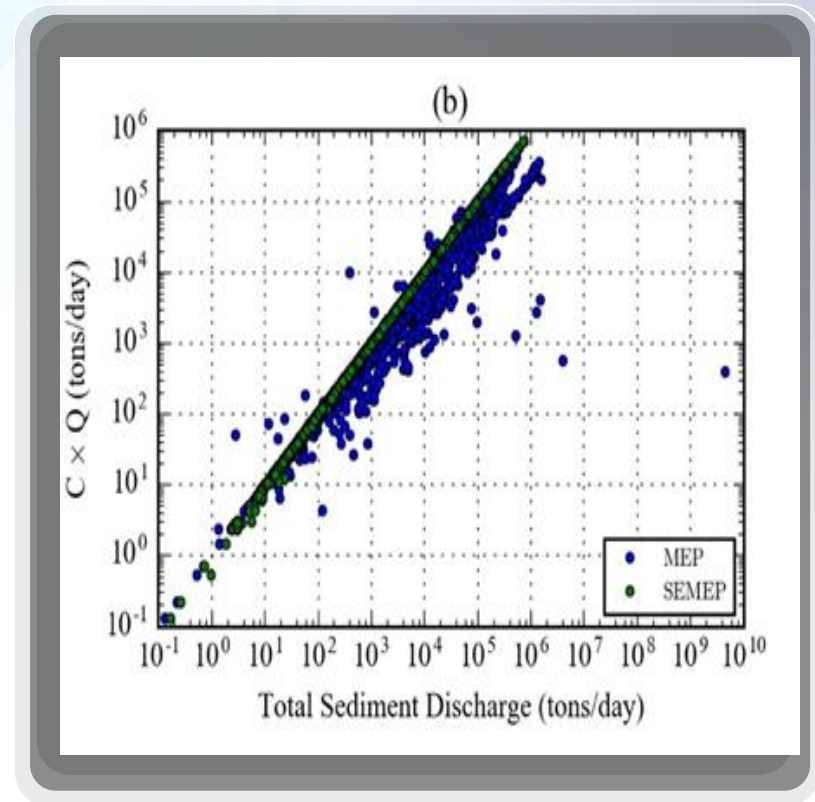
Total Sediment Discharge

SEMEP vs MEP testing in South Korea

35 Gaging stations in South Korea



MEP vs SEMEP



Sedimentation Engineering

1. Incipient motion
2. SEMEP and SEMEPP
3. **Mudflows and Debris Flows**



Physical Properties

Table 10.3. *Physical properties of mudflows and debris flows*

C_v	C_w	C_{ppm}	$C_{mg/l}$ mg/l	ρ_m^a kg/m ³	ρ_{md} kg/m ³	p_o
0.05	0.122	122,401	132,500	1,082	132	0.95
0.10	0.227	227,468	265,000	1,165	265	0.9
0.20	0.4	398,496	530,000	1,330	530	0.8
0.30	0.53	531,773	795,000	1,495	795	0.7
0.40	0.64	638,554	1,060,000	1,660	1,060	0.6
0.50	0.73	726,027	1,325,000	1,825	1,325	0.5
0.60	0.8	798,994	1,590,000	1,990	1,590	0.4
0.70	0.86	860,789	1,855,000	2,155	1,855	0.3

^a Calculations assuming $G = \rho_s/\rho = 2.65$

1. Rheology

Classification and Rheology



Total shear stress :

$$\tau = \tau_y + \tau_v + \tau_t + \tau_d$$

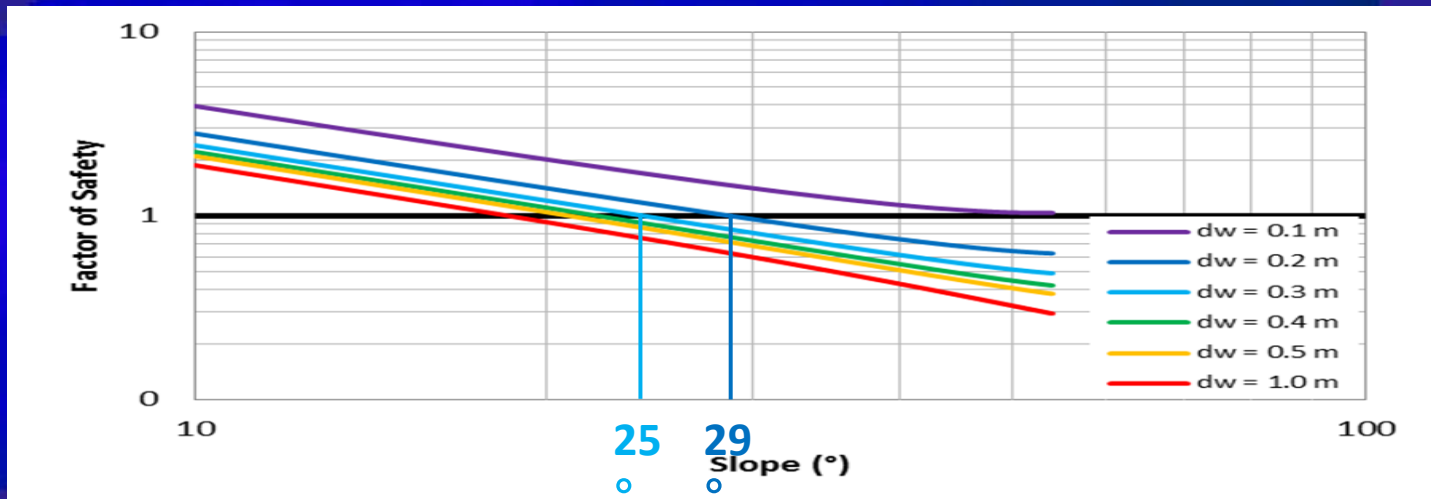
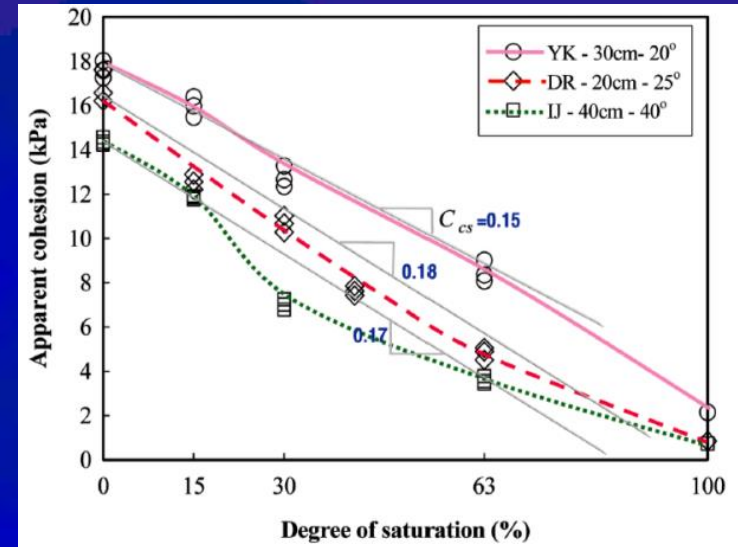
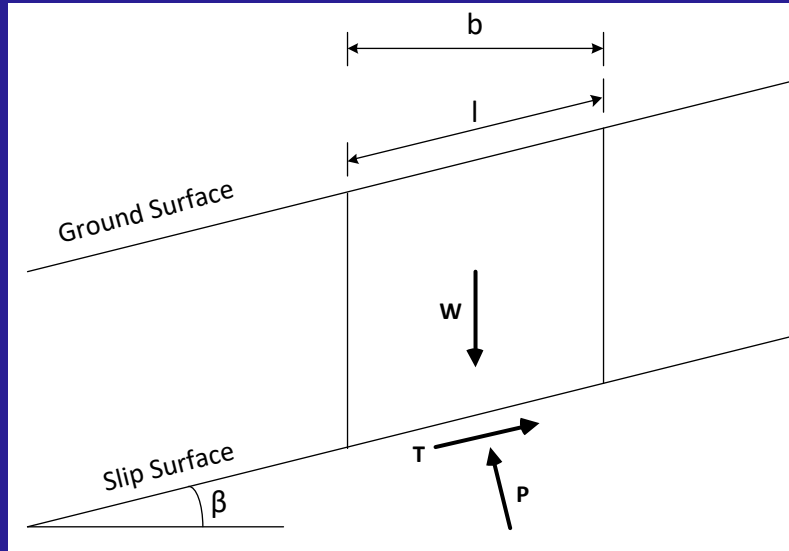
Yield stress

Viscous stress

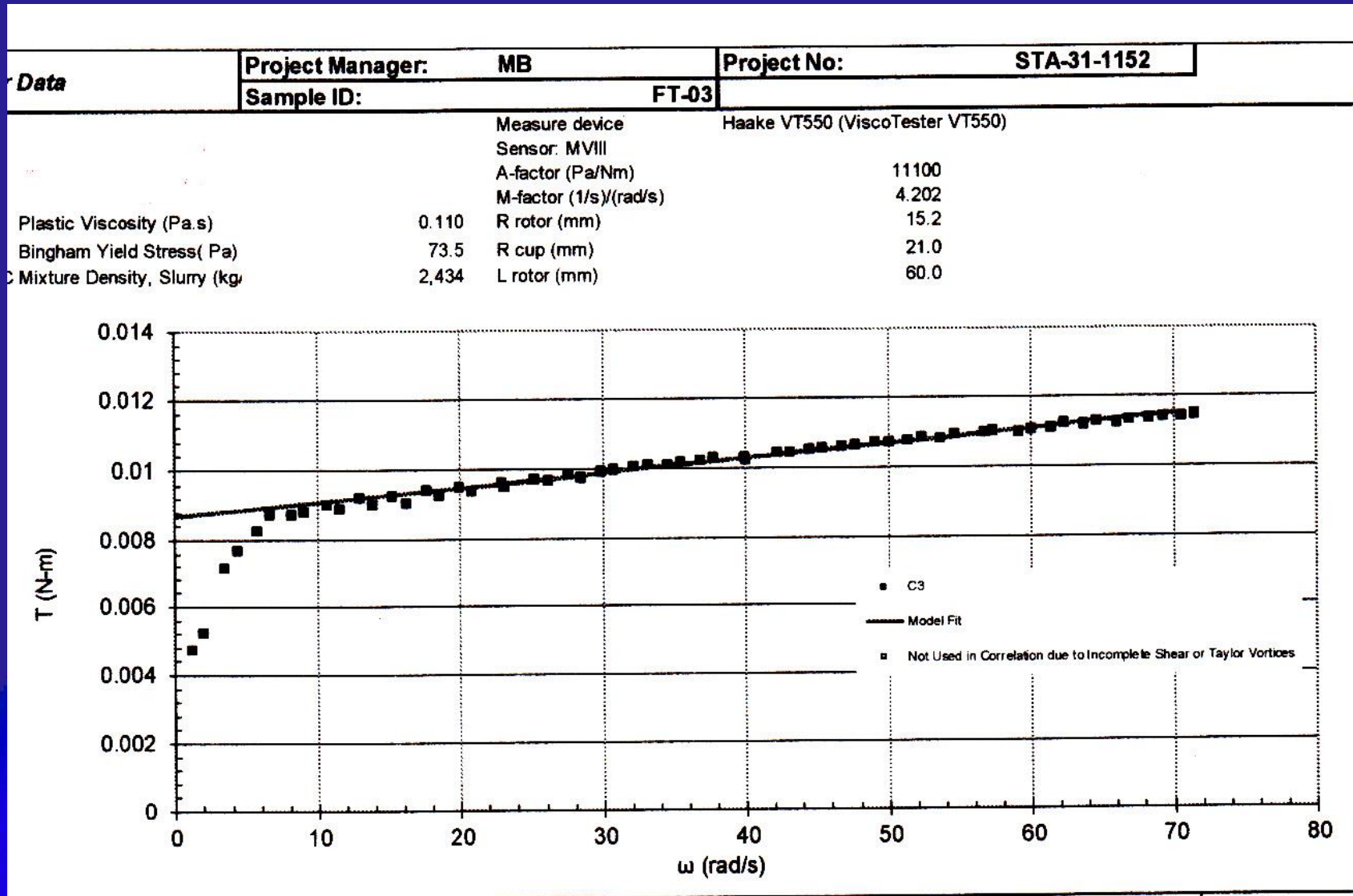
Turbulent stress

Dispersive stress

Yield Strength for Landslides



Viscometer Test



Rheogram

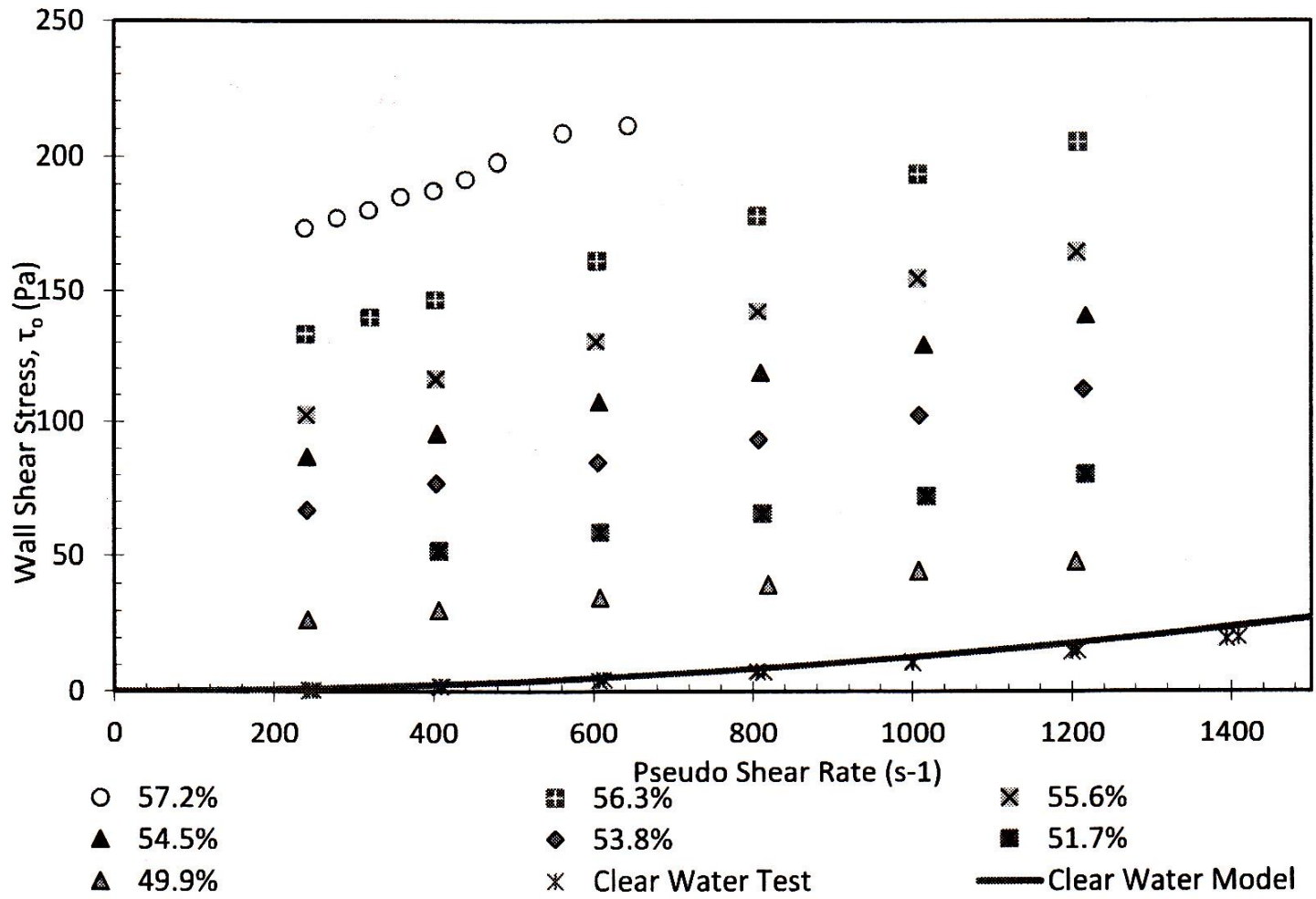


Figure 11: FT-02 Pseudo-Shear Diagram

Yield and Viscosity of Mudflows

Mudflow Rheology

$$\tau = \tau_y + \mu_m \frac{dv}{dy}$$

τ_y is the yield stress

μ_m is the dynamic viscosity (water at 0.001 Pa.s)

dv/dy is the velocity gradient (typically $1-10 \text{ s}^{-1}$)

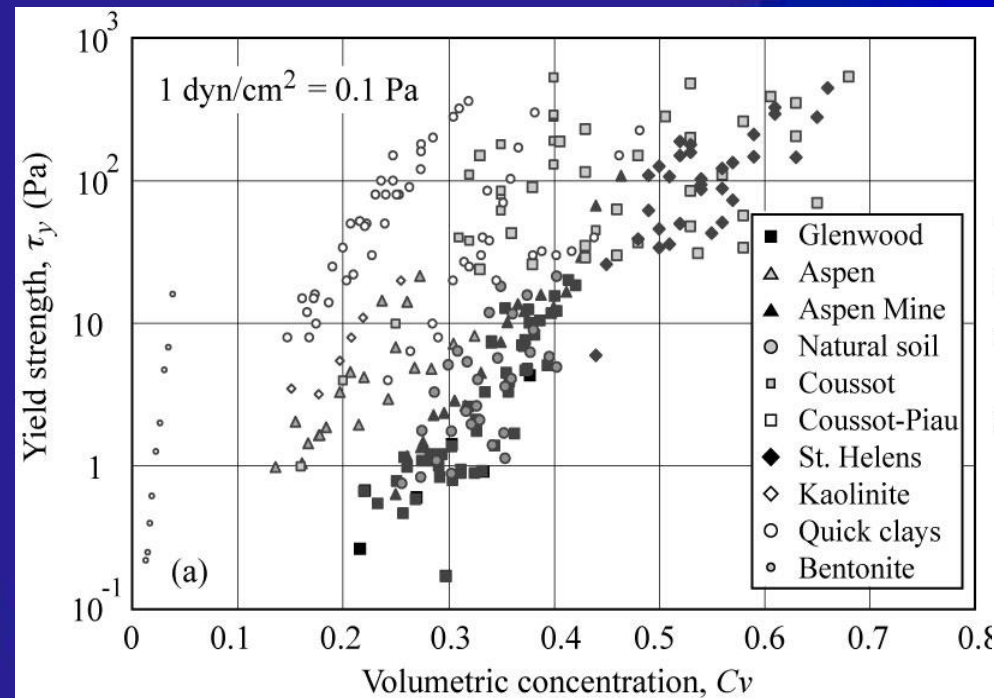


Figure 10.13a. Yield strength versus volumetric concentration

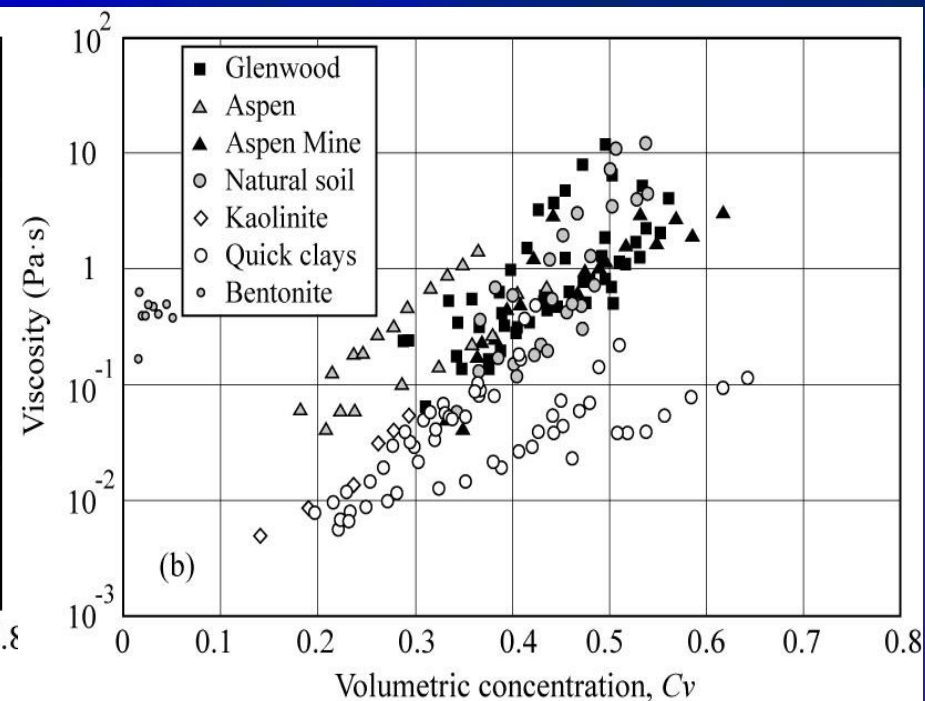


Figure 10.13b. Dynamic viscosity versus volumetric concentration

Veja o momento exato do rompimento da barragem da Vale
Jornal Hoje

25/01/2019 12:23:32



ST-01

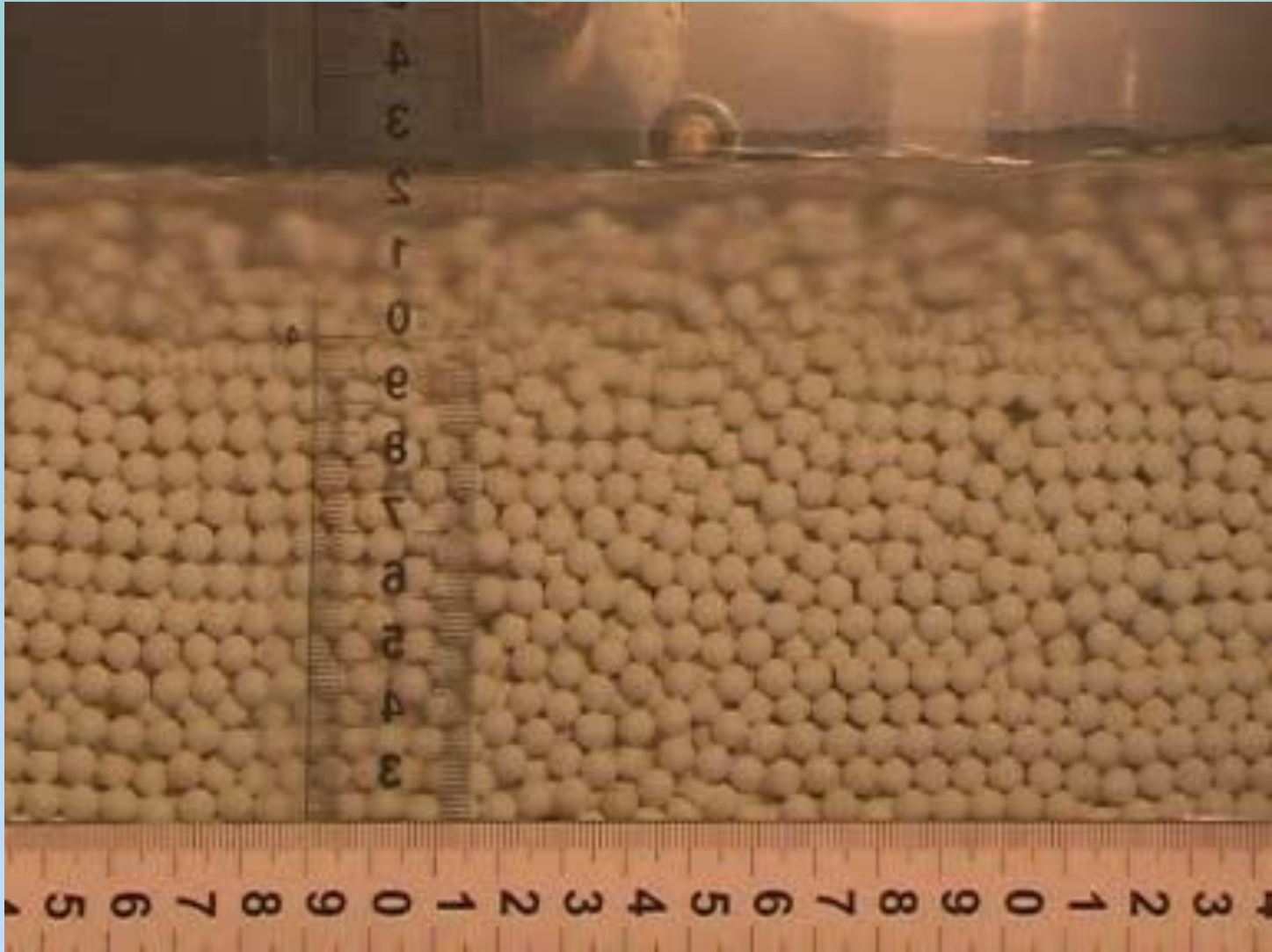
00:13 / 02:17



Turbulence



VIDEO FOOTAGE from Anna Paris and Aronne Armanini (U. Trento)



Rheology of Debris Flows

Debris flow Rheology

$$\tau = \tau_y + \mu_m \frac{dv}{dy} + \zeta \left(\frac{dv}{dy} \right)^2 \quad \text{quadratic model}$$

$$\zeta = \rho_m l_m^2 + c_{Bd} \rho_s \lambda^2 d_s^2 \quad \text{turbulent-dispersive parameter}$$

where

ρ_m is the mass density of the hyper-concentration

l_m is the mixing length, ($l_m = \kappa y$ from elevation y and von Karman constant $\kappa \cong 0.4$). Note that the mixing length is scaled to the flow depth

c_{Bd} is the Bagnold coefficient ($c_{Bd} \cong 0.01$)

ρ_s is the mass density of solids

$$\lambda = \left(\left(\frac{0.615}{C_v} \right)^{1/3} - 1 \right)^{-1} \quad \text{is Bagnold's linear concentration given the volumetric concentration } C_v, \text{ and}$$

d_s is the grain size.

The dispersive term requires three condition: (1) very high concentration, i.e. $C_v > 0.5$; (2) very large sediment where h is the flow depth; and (3) high rates of deformation, i.e. large dv/dy .

As a result, the turbulent term is dominant unless $ds > h/20$.

Turbulent or Dispersive?

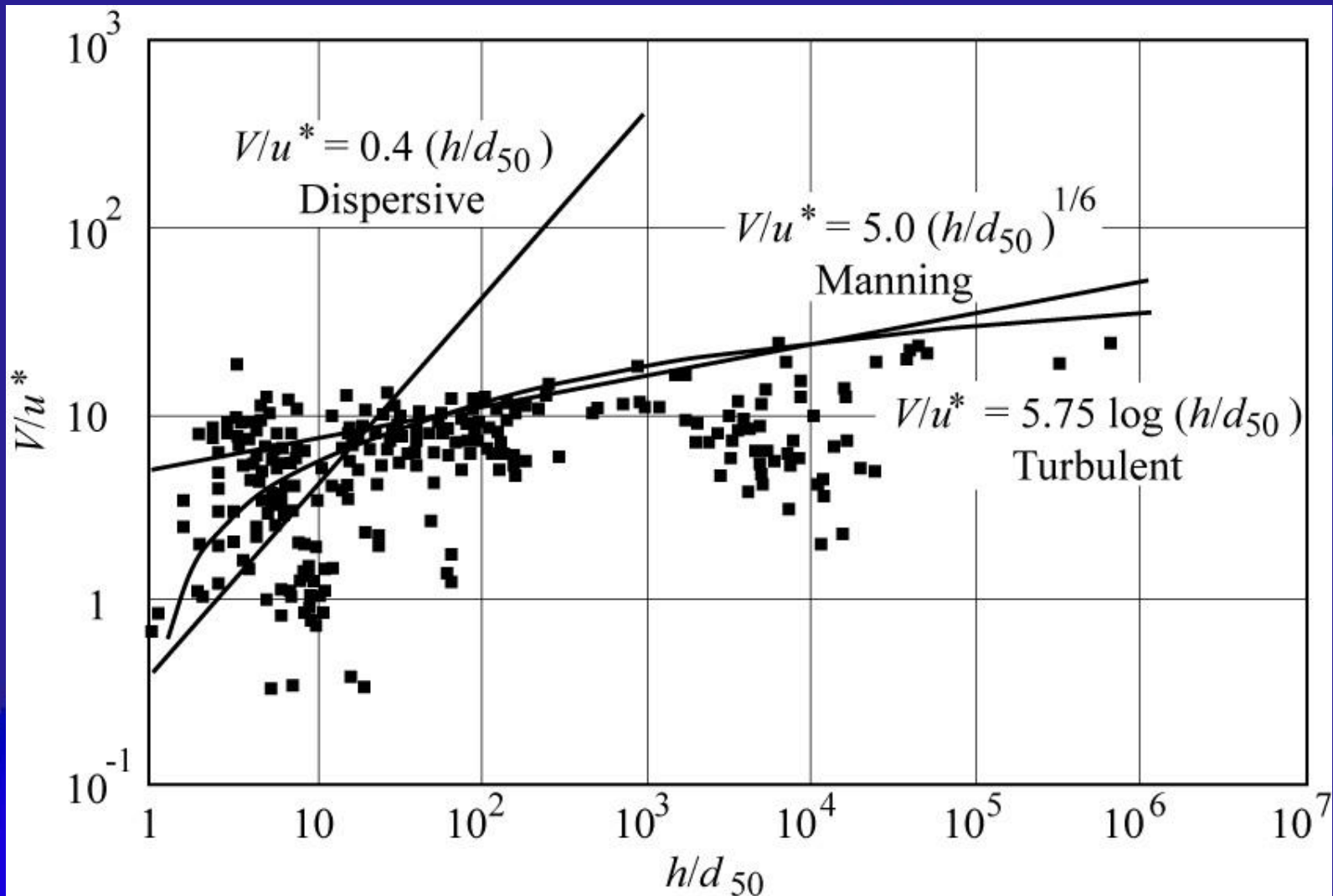


Figure 10.16. Resistance diagram for hyperconcentrated flows.

Mitigation Countermeasures

Guidelines for designing mitigation countermeasures based on the type of hyperconcentrated flow

2. Landslides

Yield Stress -> Landslides

Total shear stress :

$$\tau \sim \tau_y$$

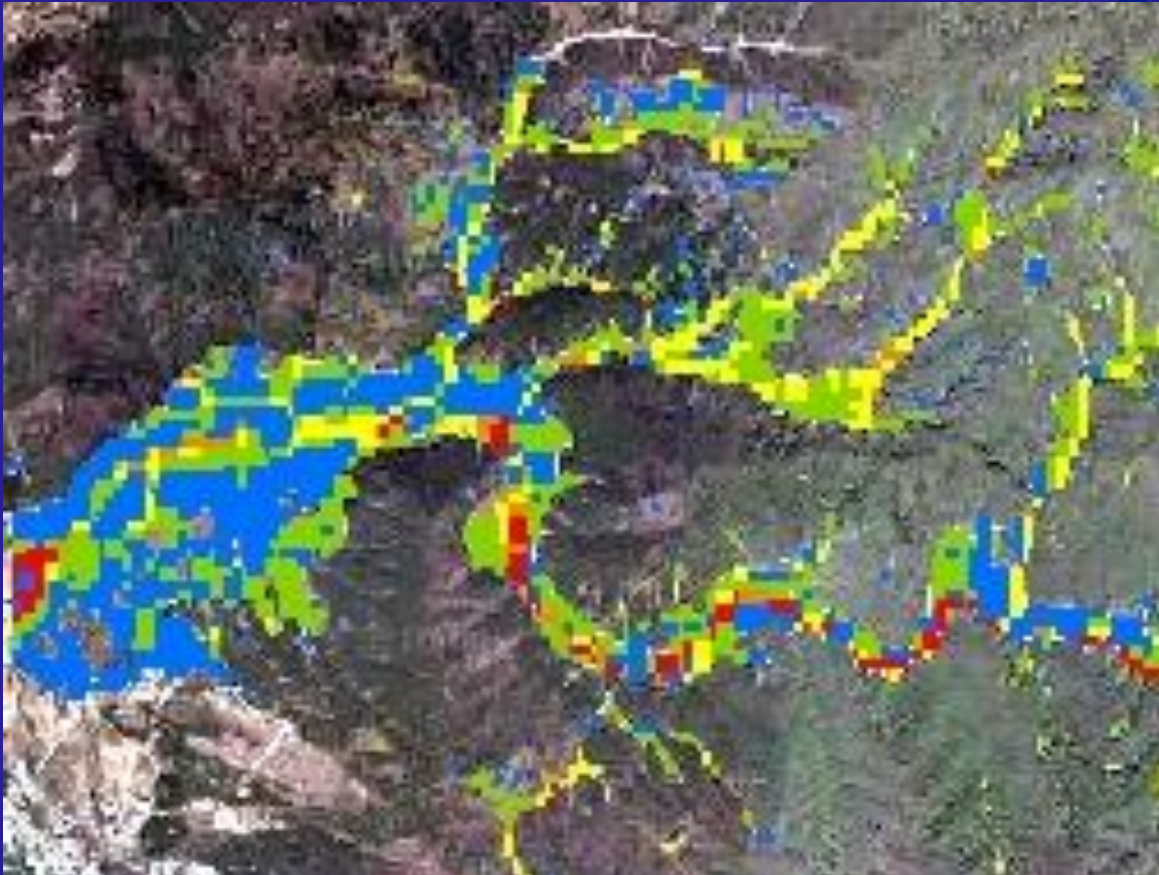
Landslides \longleftrightarrow Yield stress

Mangun mountain, South Korea



DB.CHOSUN.COM

Landslides




- Steep hillslopes
- High rainfall precipitation
- High Infiltration
- Saturated yield strength $\sim 1\text{kPa}$

Countermeasures

Landslide Features:

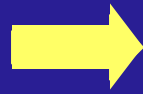
- High precipitations and infiltration
- Steep terrain
- Saturated yield strength $\sim 1\text{kPa}$

Effective Solution  Drain, Slope
Reduction, Vegetation

- Terraces
- Improved drainage

Landslide Countermeasures

**Effective
Solution**



**Slope reduction,
drain, vegetation**

- **Terraces**
- **Drainage**



3. Mudflows

Yield + Viscous -> Mudflows

Total shear stress :

$$\tau = \tau_y + \tau_v$$

Mudflow \leftrightarrow Yield stress \downarrow
Viscous stress



Mudflow




- High viscosity and yield stress
- High concentration of silts and clays
- $45\% < C_v < 55\%$
- Low velocity
- Low Froude Number
- No abrasion
- Large flow depths
- High pressure

Countermeasures

Mudflow Features:

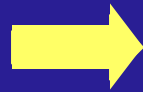
- High viscosity and yield stress
- High concentration of silt and clay
- $45\% < C_v < 55\%$
- Low Froude Number
- No abrasion

Effective Solution  Store, Deflect, Spread

- Storage basins
- Deflection walls

Mudflow Countermeasures

**Effective
Solution**



Store, Deflect, Spread

- **Storage basins**
- **Deflection walls**

Storage Basin



Deflection Wall



4. Mudfloods

Turbulence -> Mudfloods

Total shear stress :

$$\tau = \tau_y + \tau_v + \tau_t$$

Yield stress

Viscous stress

Mudfloods \longleftrightarrow Turbulent stress



Mud Flood

- Turbulent
- Non-cohesive particles
- Small particles
- C_v as high as 40%
- High velocity
- High Froude Number
- Abrasive

Straight Channel



Countermeasures

Mud Flood Features:

- Turbulent
- Non cohesive particles
- Cv as high as 40%
- High Froude Number
- Abrasive

Effective Solution  Convey

- Straight channels
- Lined canals, berm and levee channels
- Drop structures, energy dissipators

Mudflood Countermeasures

Effective Solution  Increased conveyance

- Straight channel
- Lined canal
- Berm and levee
- Drop structure

Straight Channel



Lined canal with drop structures



5. Debris Flows

Dispersive -> Debris Flows

Total shear stress :

$$\tau = \tau_y + \tau_v + \tau_t + \tau_d$$

Landslides ↔ Yield stress

Mudflows ↔ Viscous stress

Mudflows ↔ Turbulent stress

Debris flows ↔ Dispersive stress



Los Corales



Los Corales



Los Corales



Debris Flow




- Dispersive
- Large clastic particles
- Non cohesive
- Low viscosity
- High velocity
- Destructive impact force

Countermeasures

Debris Flow Features:

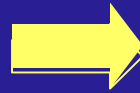
- Dispersive
- Large clastic particles
- Low viscosity
- Large velocity
- High impact

Effective Solution  Retain large clasts
Drain fluid matrix

- Concrete Sabo dams
- Steel frames and debris rakes

Debris flow Countermeasures

Effective Solution



Retain large rocks
Drain water

- Concrete sabo dams
- Steel Frames
- Debris Racks

Sabo Dam and Steel Frames



Sabo Dam Construction

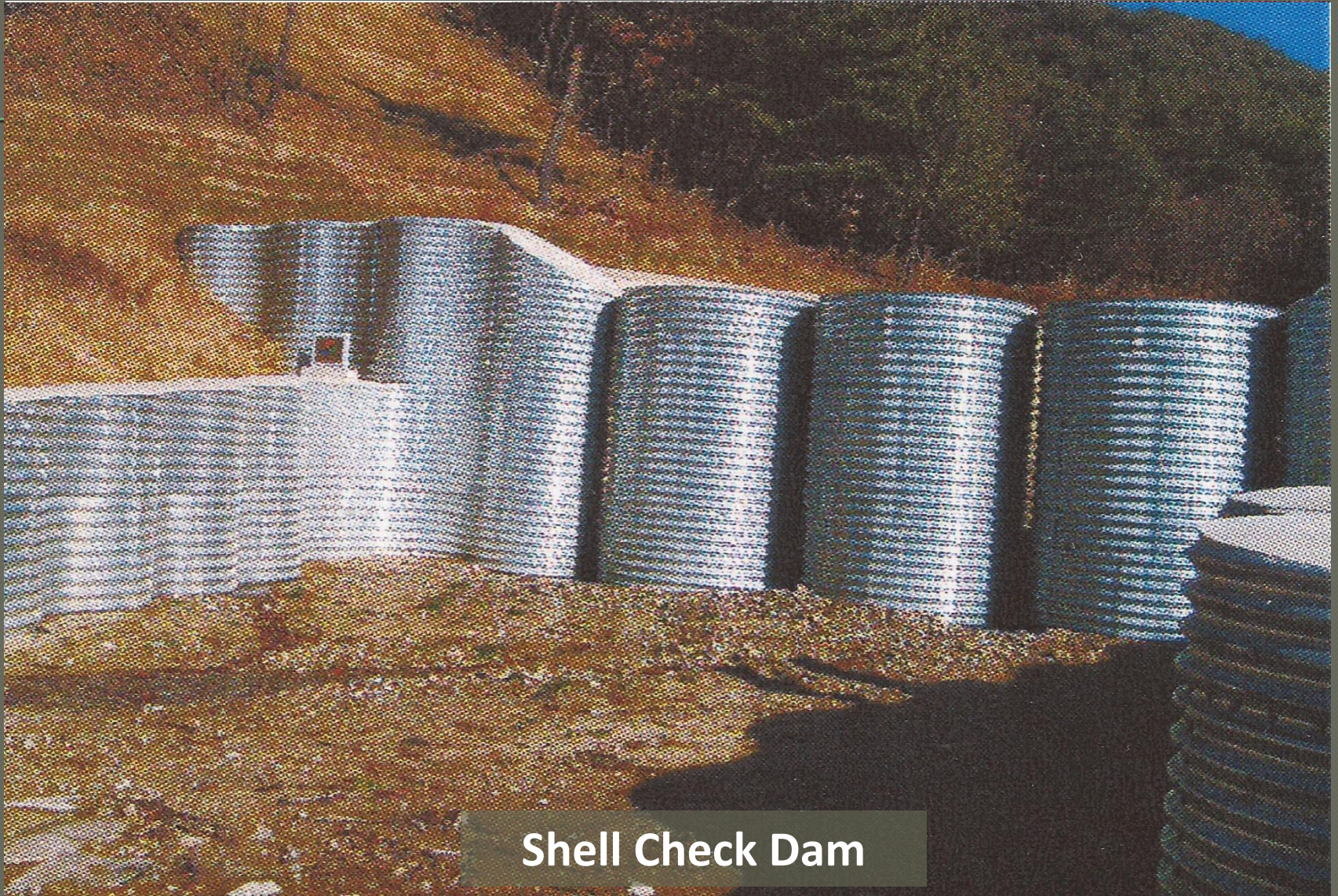


Debris Rakes



Sabo Dam and Steel Frames





Shell Check Dam

Classification and Rheology

Total shear stress :

$$\tau = \tau_y + \tau_v + \tau_t + \tau_d$$

Landslides ↔ Yield stress

Mudflows ↔ Viscous stress

Mudflows ↔ Turbulent stress

Debris flows ↔ Dispersive stress

Conclusions

- Mitigation structures for mudflows
 - » Detention basins
 - » Deflection walls
- Mitigation structures for mud floods
 - » Straight channels
 - » Lined canals, berm and levee channels
 - » Drop structures, energy dissipators
- Mitigation structures for debris flows
 - » Concrete Sabo dams
 - » Steel frames and debris rakes

6. Example (Quiz?)

Mt Umyeon

특집 뉴스라인
집중호우 강타

RBS100









Summary and Conclusions

1. Incipient Motion

Incipient motion depends on ratio of moments, not forces.

2. SEMEP and SEMEPP

The Einstein integrals are accurately solved numerically.

3. Mudflows and Debris Flows

The flow type depends on flow rheology.

Effective countermeasures depend on the rheology.



PIERRE Y. JULIEN

Erosion and Sedimentation

Second Edition



CAMBRIDGE

Thank You!

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